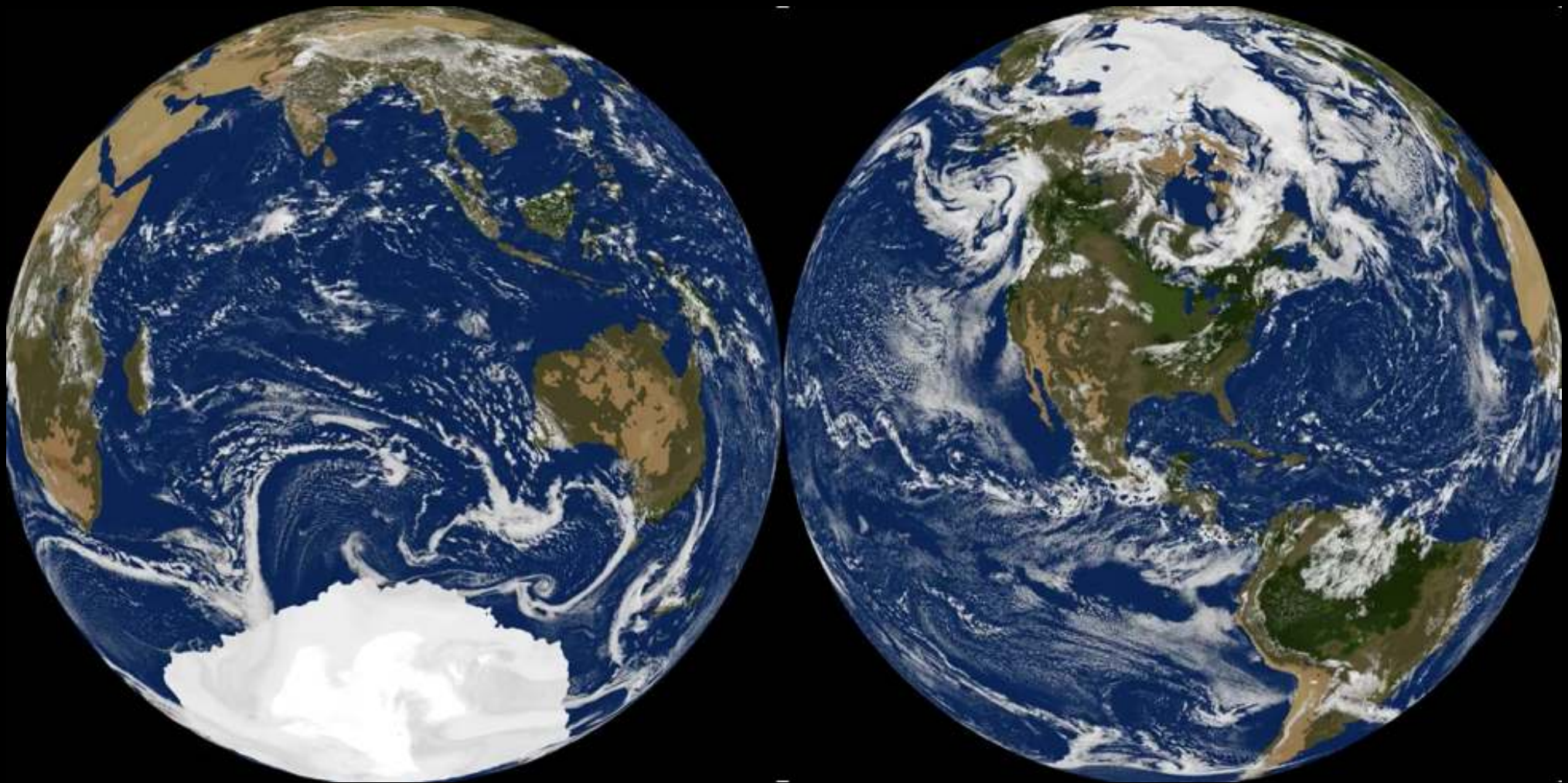


Global SAM

Marat Khairoutdinov



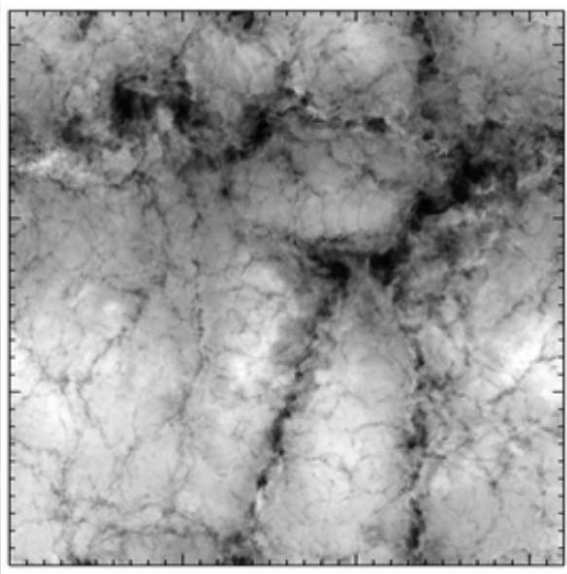
Stony Brook University
Long Island, New York, USA

DYAMOND Workshop, Mainz, Germany June 19–20, 2019

SAM: System for Atmospheric Modeling

Cloud-Resolving Model (CRM) / Large-Eddy Simulation (LES) Model

- Has been around since late 90s
- Anelastic (no sound waves), non-hydrostatic; 2D or 3D;
- Smagorinski 1st order or 1.5 prognostic TKE SGS closure;
- Radiation packages: NCAR CAM3 and RRTM;
- 2nd-order space and 3rd-order Adams-Bashforth time differences for momentum;
- Conservation of momentum and kinetic energy;
- fully 3-D positive definite and monotonic transport (MPDATA) for all scalars;
- Several cloud microphysics packages of various complexity;
- Comprehensive land-surface model with different types of vegetation and interactive soil
- "Box-fill" method for the topography;
- Massively parallel (domain decomposition, MPI);



Spherical coordinates (Lat-Lon grid) : φ - latitude, $\mu = \cos(\varphi)$

$x = r\lambda$, $y = r\varphi$, where r is Earth radius.

Cartesian grid when $\mu = 1$ ($r = \infty$)

Momentum (anelastic approximation, no sound waves):

$$\frac{\partial u}{\partial t} + \frac{1}{\mu} \frac{\partial}{\partial x} uu + \frac{1}{\mu \rho_0} \frac{\partial}{\partial y} \mu \rho_0 uv + \frac{1}{\rho_0} \frac{\partial}{\partial z} \rho_0 uw = -\frac{1}{\mu} \frac{\partial \pi}{\partial x} + \left(f + \frac{u}{r} \tan \varphi \right) v + Du$$

$$\frac{\partial v}{\partial t} + \frac{1}{\mu} \frac{\partial}{\partial x} vu + \frac{1}{\mu \rho_0} \frac{\partial}{\partial y} \mu \rho_0 vv + \frac{1}{\rho_0} \frac{\partial}{\partial z} \rho_0 vw = -\frac{\partial \pi}{\partial y} - \left(f + \frac{u}{r} \tan \varphi \right) u + Dv$$

$$\frac{\partial w}{\partial t} + \frac{1}{\mu} \frac{\partial}{\partial x} wu + \frac{1}{\mu \rho_0} \frac{\partial}{\partial y} \mu \rho_0 wv + \frac{1}{\rho_0} \frac{\partial}{\partial z} \rho_0 ww = -\frac{\partial \pi}{\partial z} + B + Dw$$

Continuity:

$$\frac{1}{\mu} \frac{\partial}{\partial x} u + \frac{1}{\mu \rho_0} \frac{\partial}{\partial y} \mu \rho_0 v + \frac{1}{\rho_0} \frac{\partial}{\partial z} \rho_0 w = 0$$

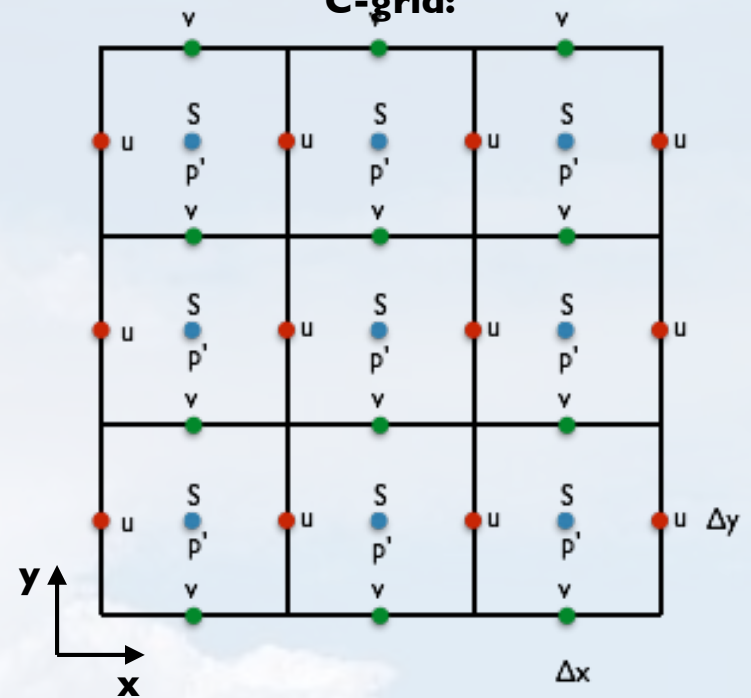
Scalars :

$$\frac{\partial S}{\partial t} + \frac{1}{\mu} \frac{\partial}{\partial x} wS + \frac{1}{\mu} \frac{\partial}{\partial y} \mu vS + \frac{1}{\rho_0} \frac{\partial}{\partial z} \rho_0 wS = Q + Dw$$

SGS Diffusion:

$$D\psi = \frac{1}{\mu^2} \frac{\partial}{\partial x} k_h \frac{\partial \psi}{\partial x} + \frac{1}{\mu \rho_0} \frac{\partial}{\partial y} \mu k_h \frac{\partial \psi}{\partial y} + \frac{1}{\rho_0} \frac{\partial}{\partial z} \rho_0 k_z \frac{\partial \psi}{\partial z}$$

C-grid:



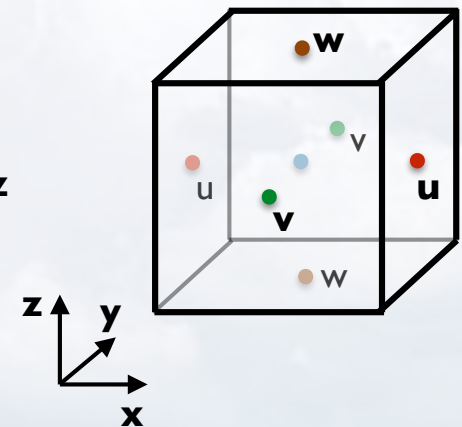
Still same horizontal grid with constant Δx and Δy

Single Reference State profiles:

$$\rho_o = \rho_o(z)$$

$$T_o = T_o(z)$$

Pressure Solver:
FFT in x , multigrid in $y-z$



Diagnostic Equation for Pressure p'

Cartesian grid

Poisson equation:

$$\frac{\partial^2 p'}{\partial x^2} + \frac{\partial^2 p'}{\partial y^2} + \frac{\partial}{\partial z} \bar{\rho} \frac{\partial p'}{\partial z} \frac{1}{\bar{\rho}} = F$$

3-D elliptic equation
on Cartesian grid

Fourier transformation in x and y (FFT):

$$p' = \sum_{mn} \hat{p}_{mn} e^{i(mx+ny)} \quad F = \sum_{mn} \hat{F}_{mn} e^{i(mx+ny)}$$

Works when Δx and Δy are constant

Diagnostic equation for Fourier coefficients:

$$-(m^2 + n^2) \hat{p}_{mn} + \frac{1}{\bar{\rho}} \frac{\partial}{\partial z} \bar{\rho} \frac{\partial \hat{p}_{mn}}{\partial z} = \hat{F}_{mn}$$

1-D equation - easy! ($N_x N_y$ of them)

Diagnostic Equation for Pressure p'

Lat-Lon Grid

Lat-Lon grid - Poisson equation:

$$\frac{\partial^2 p'}{\partial x^2} + \mu \frac{\partial}{\partial y} \mu \frac{\partial p'}{\partial y} + \mu^2 \frac{\partial}{\partial z} \bar{\rho} \frac{\partial p'}{\partial z} = F \mu^2$$

Remember: $\mu = \mu(y)$

Fourier transformation in x only (FFT):

$$p' = \sum_m \hat{p}_m e^{imx} \quad F = \sum_m \hat{F}_m e^{imx}$$

Δx is constant, but coefficients μ are functions of y - can't use FFT in y as in Cartesian!

Need to solve 2-D Helmholtz (Poisson for $m=0$) equation with Neumann boundary conditions:

$$-m^2 \hat{p}_m + \mu \frac{\partial}{\partial y} \mu \frac{\partial \hat{p}_m}{\partial y} + \mu^2 \frac{\partial}{\partial z} \bar{\rho} \frac{\partial \hat{p}_m}{\partial z} = \hat{F}_m \mu^2$$

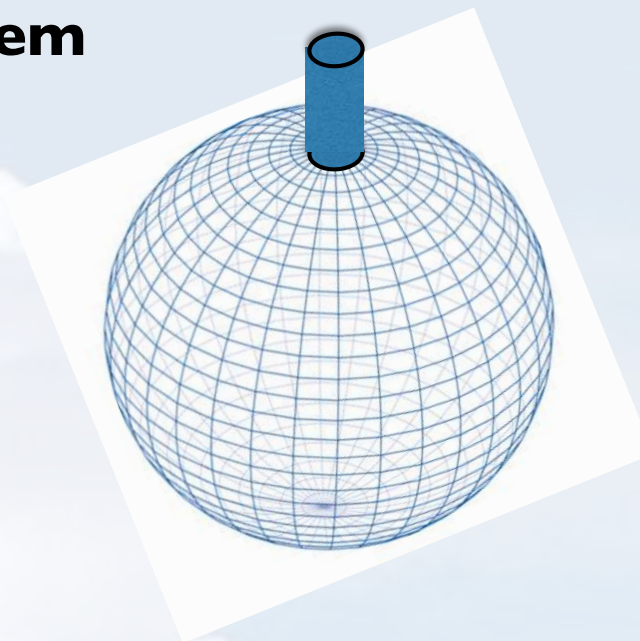
2-D (y-z) equation (N_x of them)

Solution: Custom made Geometric Multi-Grid (GMG) solver for Highly Anisotropic and Singular Elliptic problems with Neumann BCs

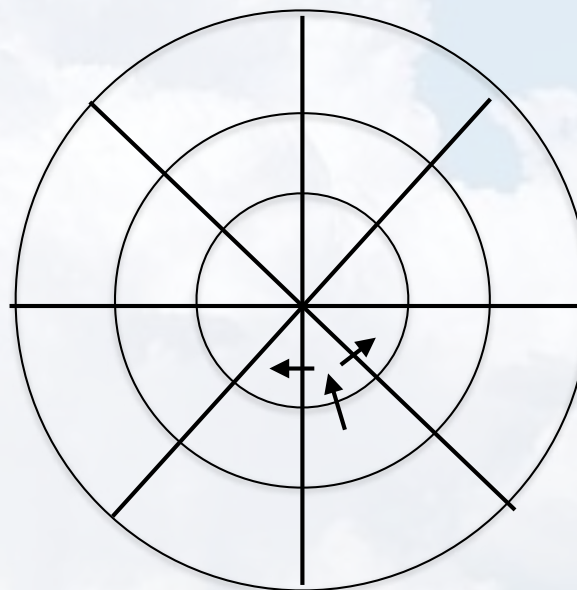
Typical convergence is a single V-Cycle even in the presence of hurricanes!

Pole Problem

**Used for DYAMOND run:
a wall around the pole at
89°N-S.**



**Current:
Triangular grid cells at the pole;
Algebraic grid coarsening**



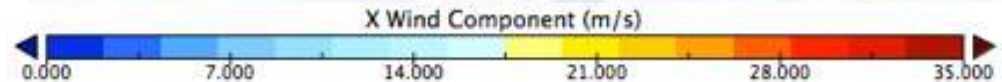
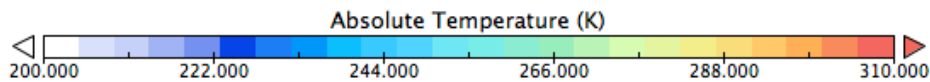
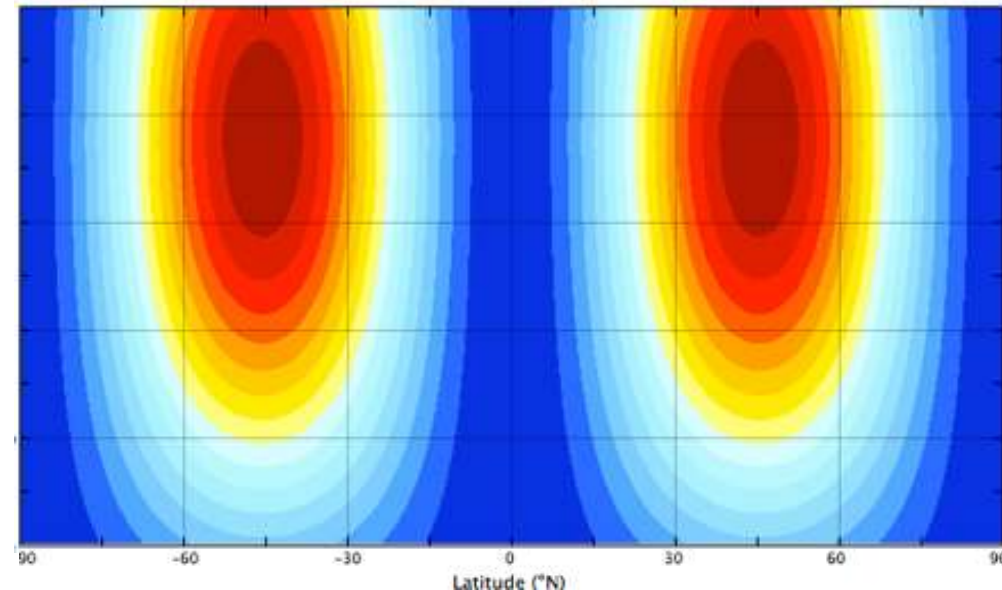
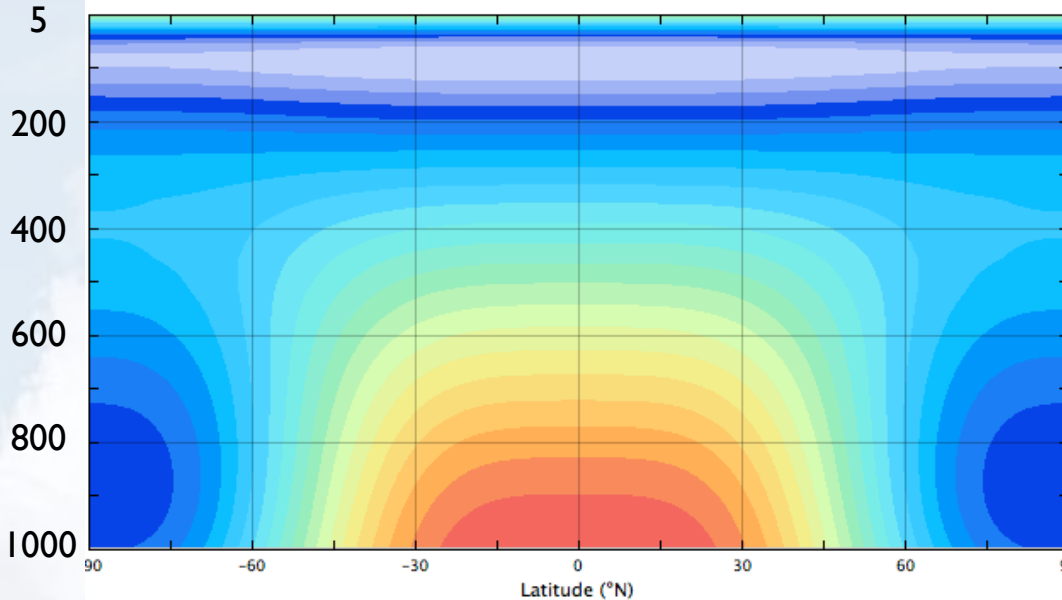
How adequate is the anelastic approximation, in particular, a single reference profile, for a global model?

Dry baroclinic-wave test (Jablonowski and Williamson 2006)

mb

Prescribed Temperature

Exact Thermal Wind solution



Setup:

- 578x288x74 grid; 0.62 x 0.62 degrees
- Single reference profiles as a global mean profiles of T and ρ ;
- Initial conditions: u from analytical solution, v=0;
- No friction, no SGS viscosity, no damping;
- Reference profiles are the global mean profiles;
- Run for 30 days;

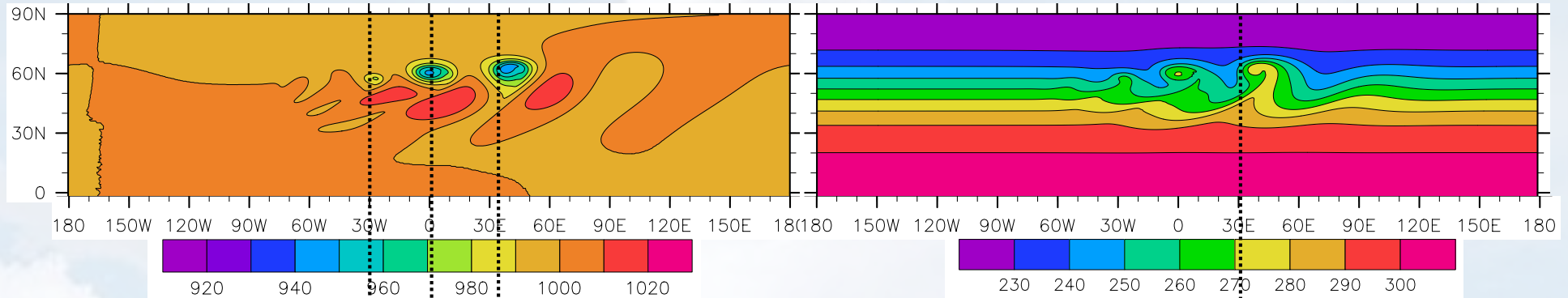
Dry Baroclinic Wave Test (Jablonowski and Williamson 2006)

SAM (0.62° x 0.62°)

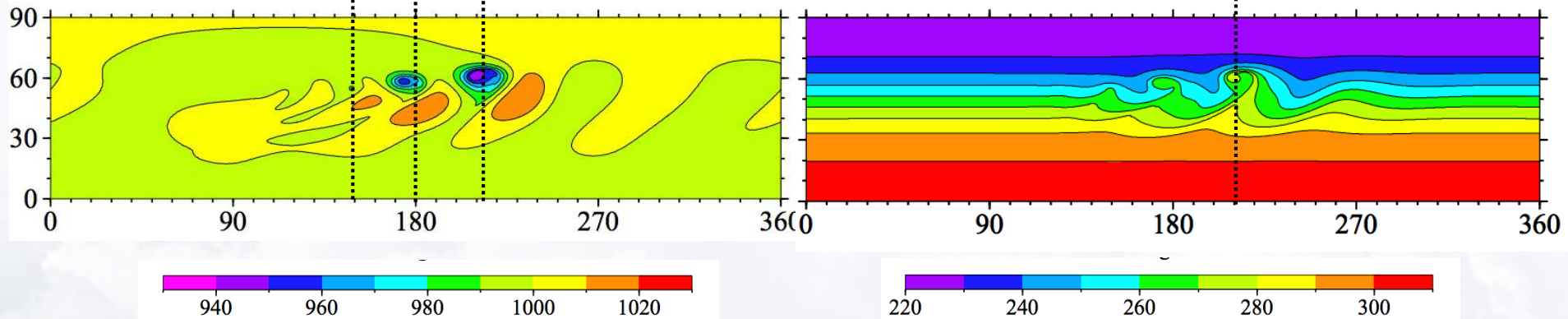
Day 9

Surface Pressure

850 hPa Temperature



CAM-FV (0.5° x 0.625°)

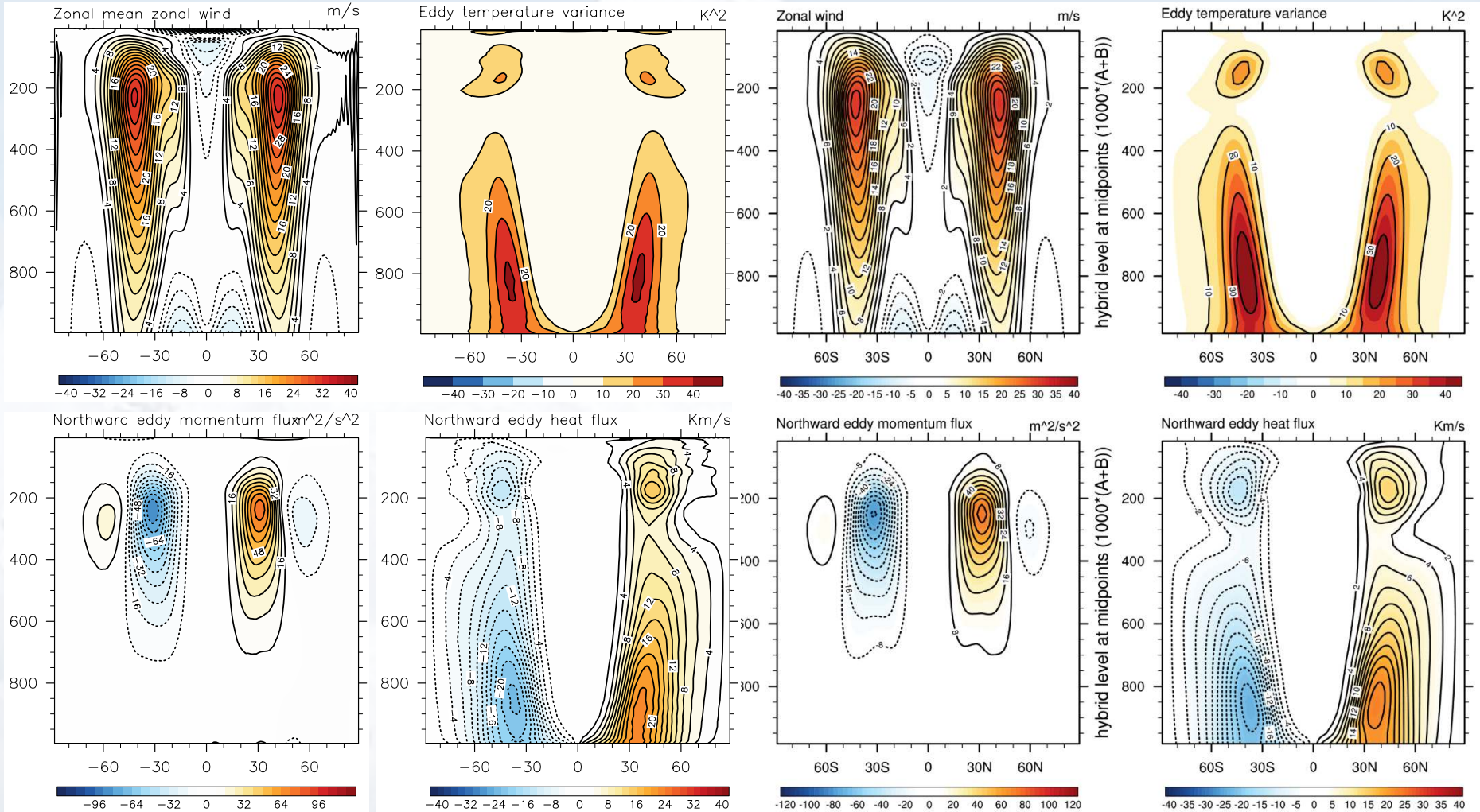


Problem: SAM's baroclinic wave propagation is a bit faster than in other global models

Held-Suarez Dry DyCore Test

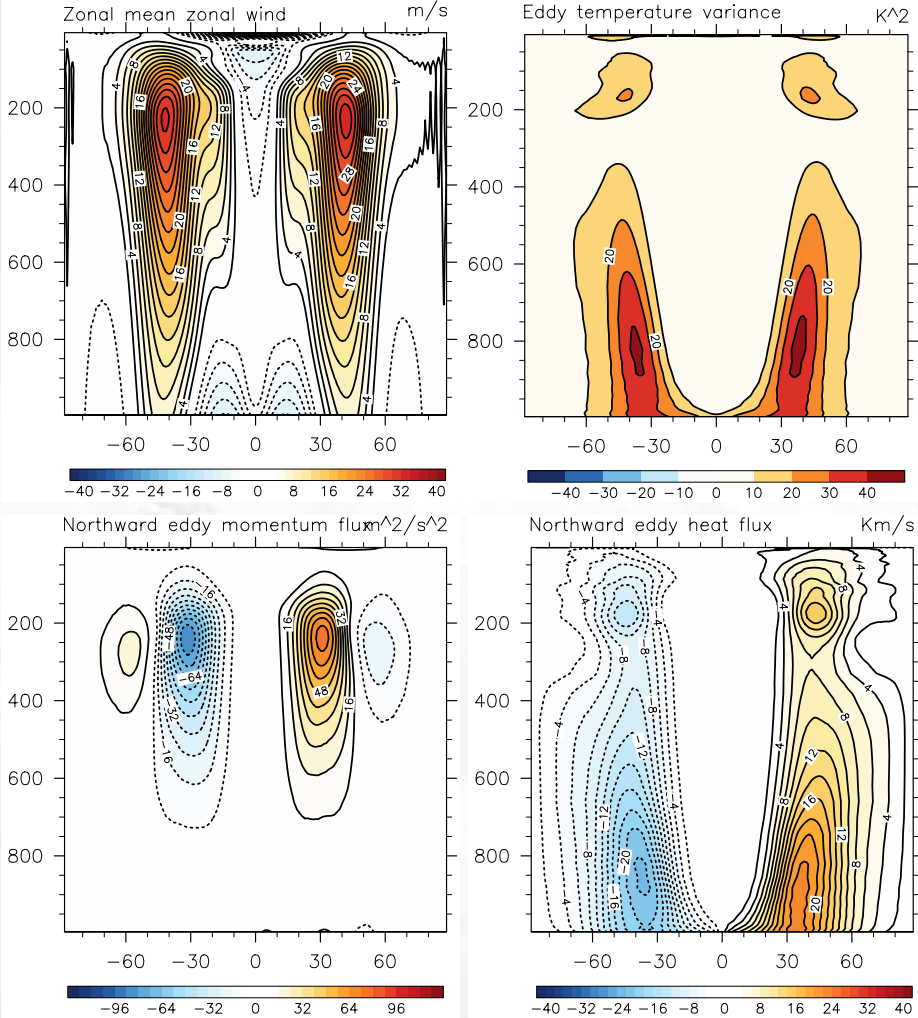
SAM (0.6° x 0.6°)

CAM-EUL (T85)

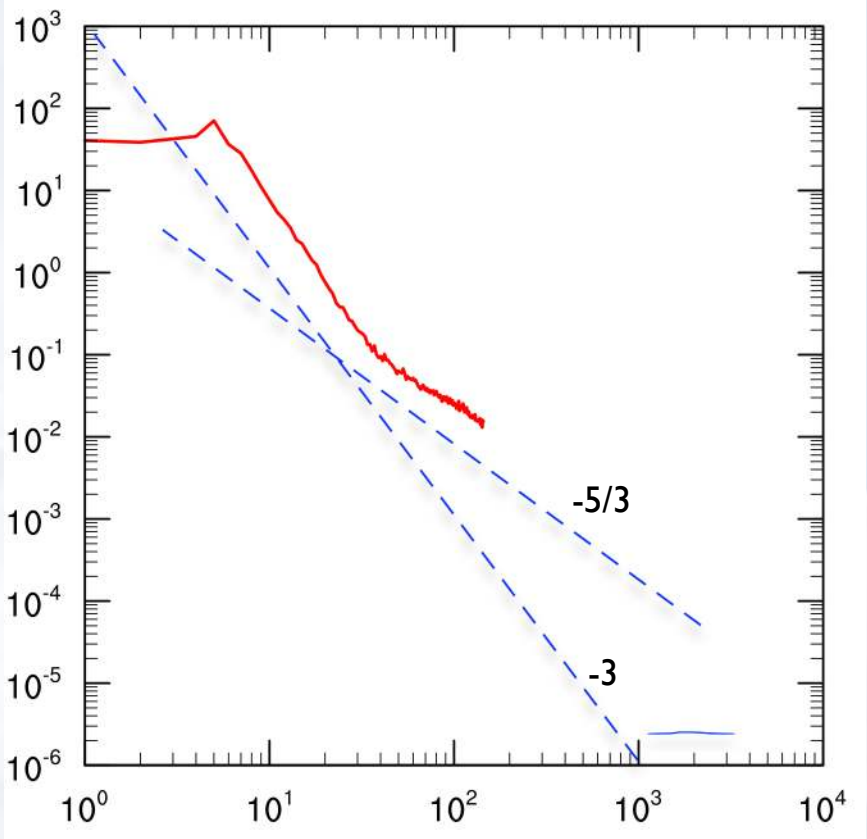


Held-Suarez Dry DyCore Test

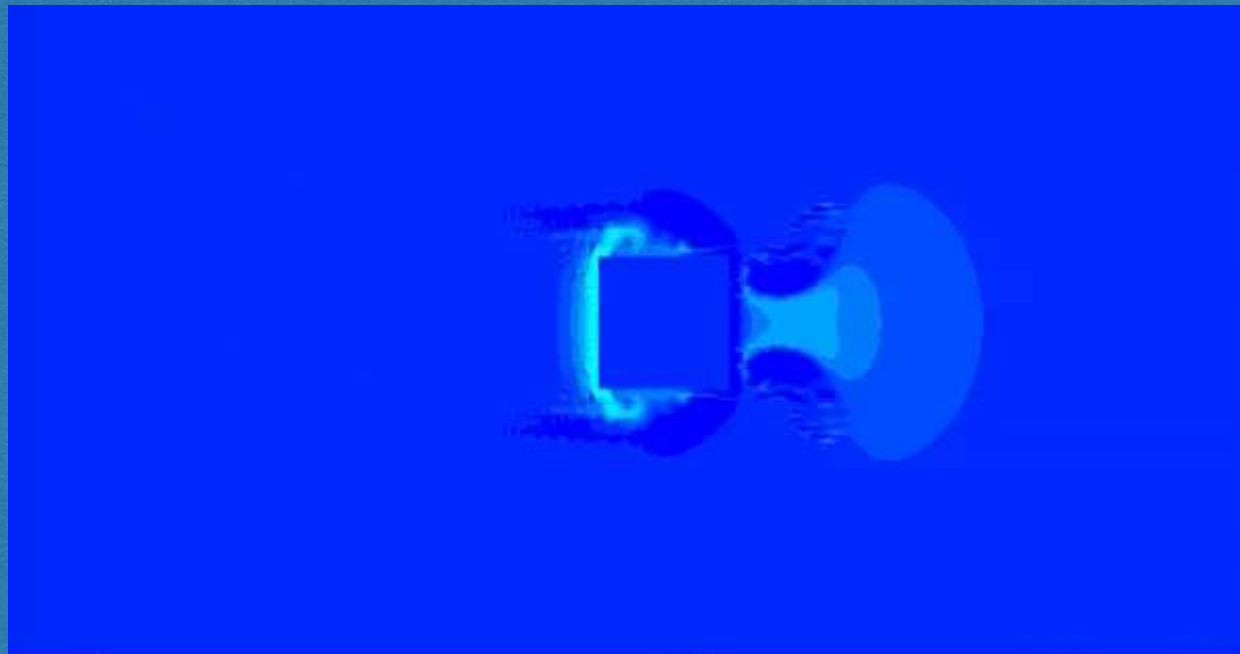
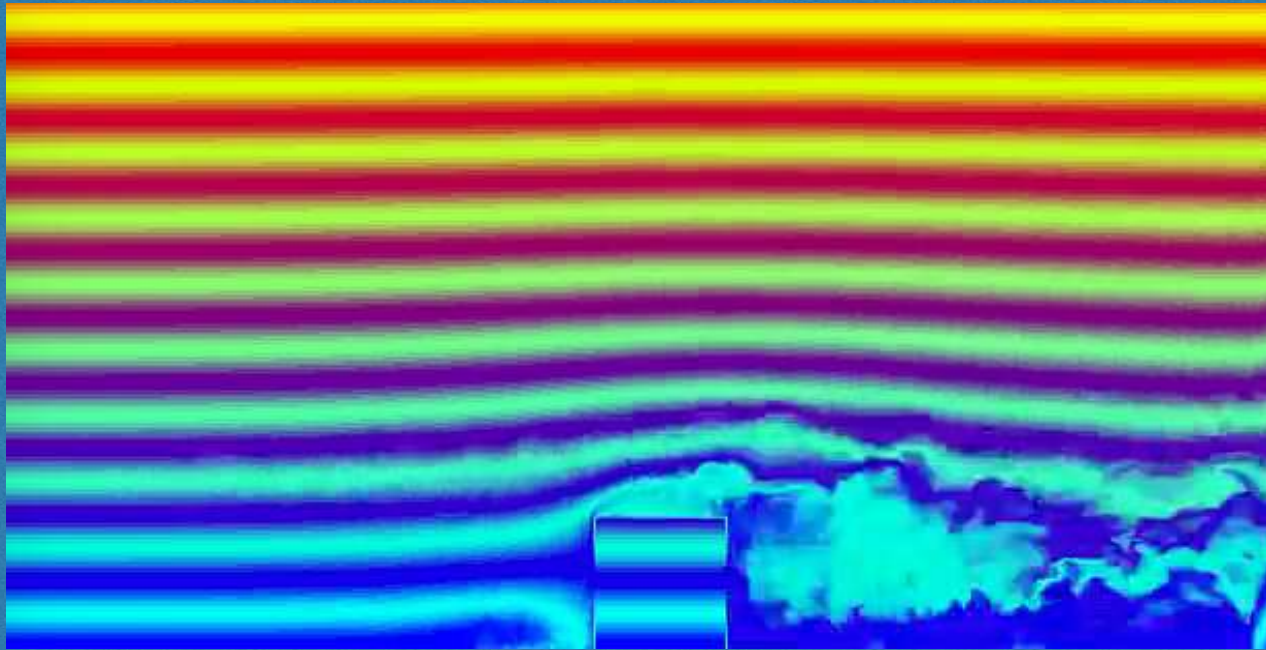
SAM (0.62° x 0.62°)



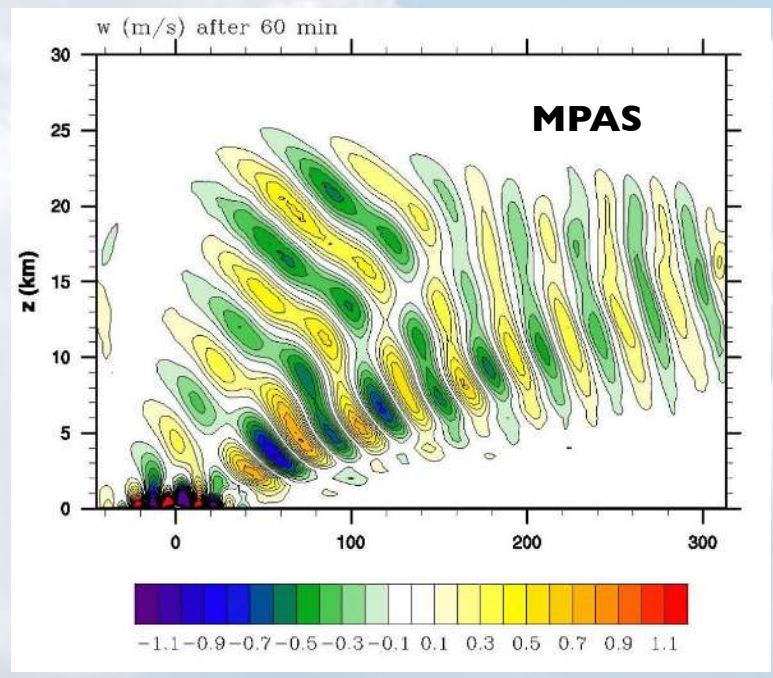
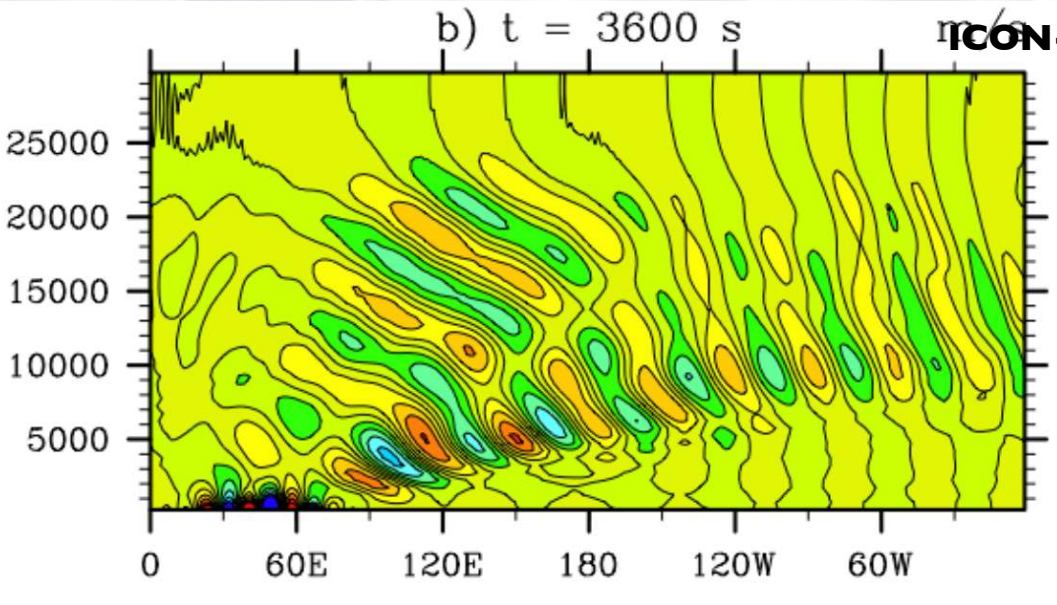
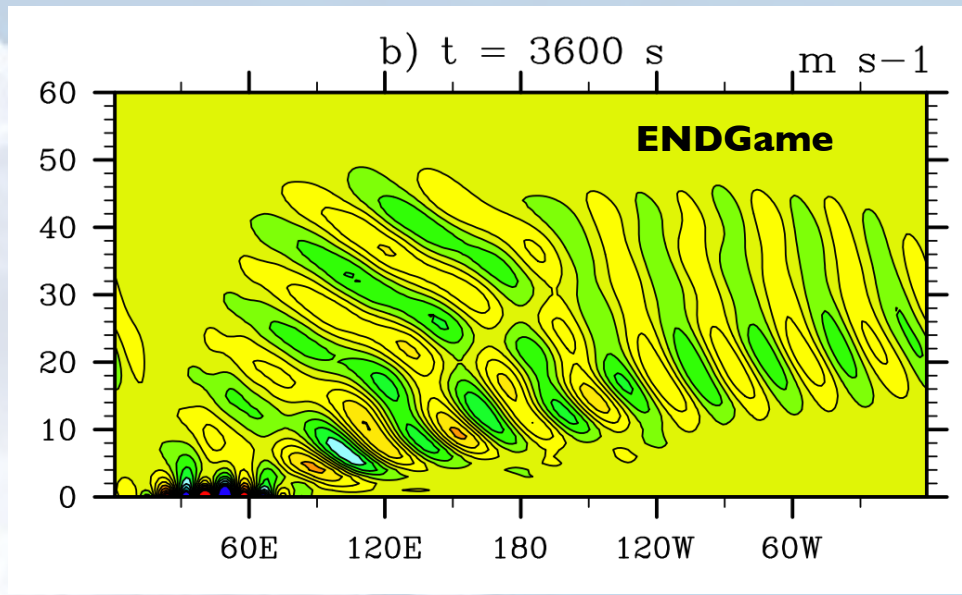
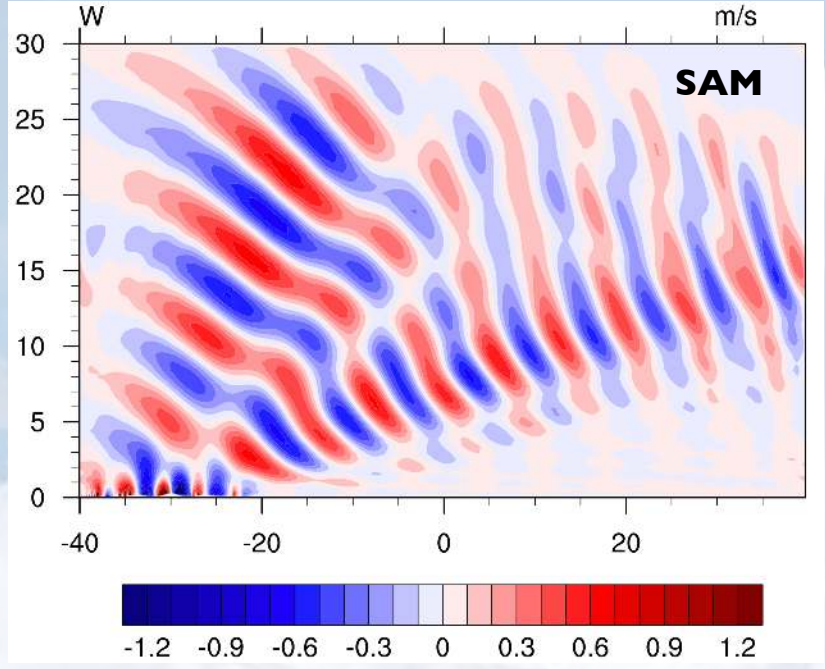
Zonal velocity spectrum mid-latitudes



Simulation of a cube in wind tunnel "Fill-box" method



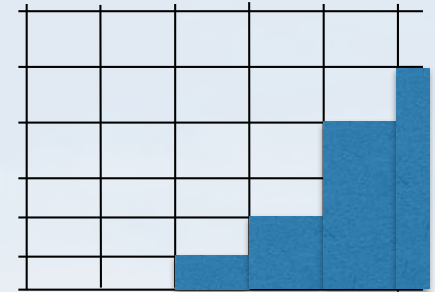
DCMIP 2.1 case: Mountain waves over a Schaar-type mountain on a small planet without shear ($X=500$): 80 levels, only 5 levels below mountain top at 250m



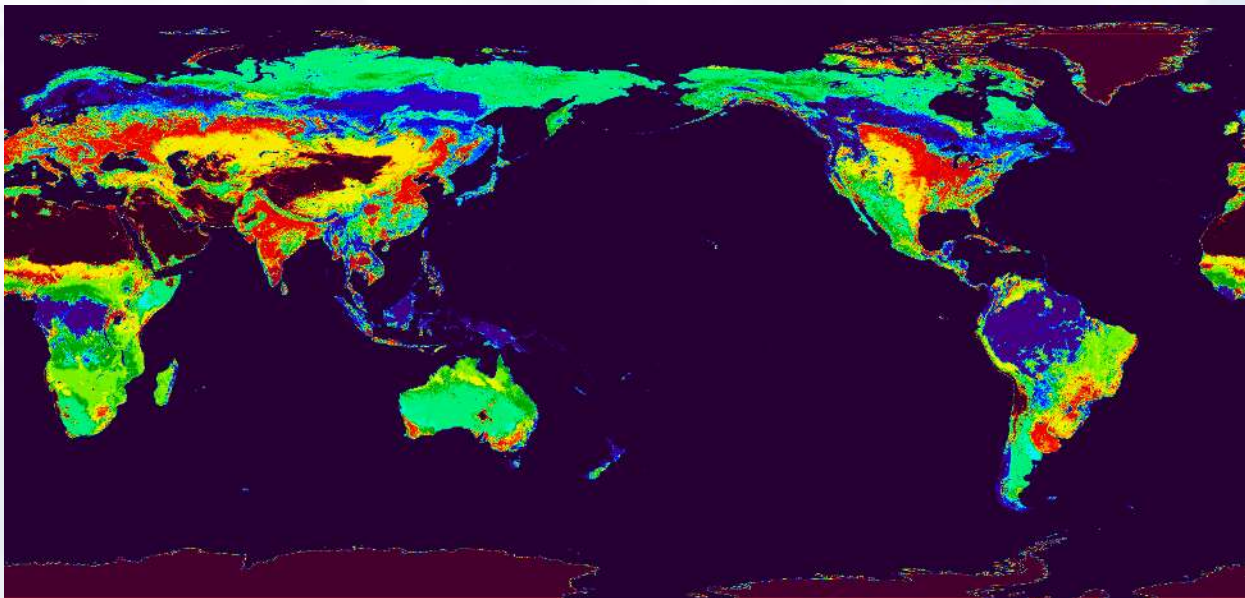
SAM Configuration for DYAMOND

- Grid 4608 x 9216 x 74 (0.039° x 0.039°, grid spacing 4.25 km at equator)
- Block topography at 4 km resolution; no smoothing except for Antarctica
- Single-Moment microphysics; CAM3 radiation called every 15 min
- Land: Simplified Land Model (Lee and Khairoutdinov 2015), 16 IGBP types, interactive snow
- Soil: 9 layers with thicknesses from 1 cm to 1 m; total depth 2.5 m
- Period: August 1st - September 10, 2016
- Performance: 6 simulated days per wall-clock day on 4608 cores (time step 7.5s)

Block-topography



16 IGBP land types

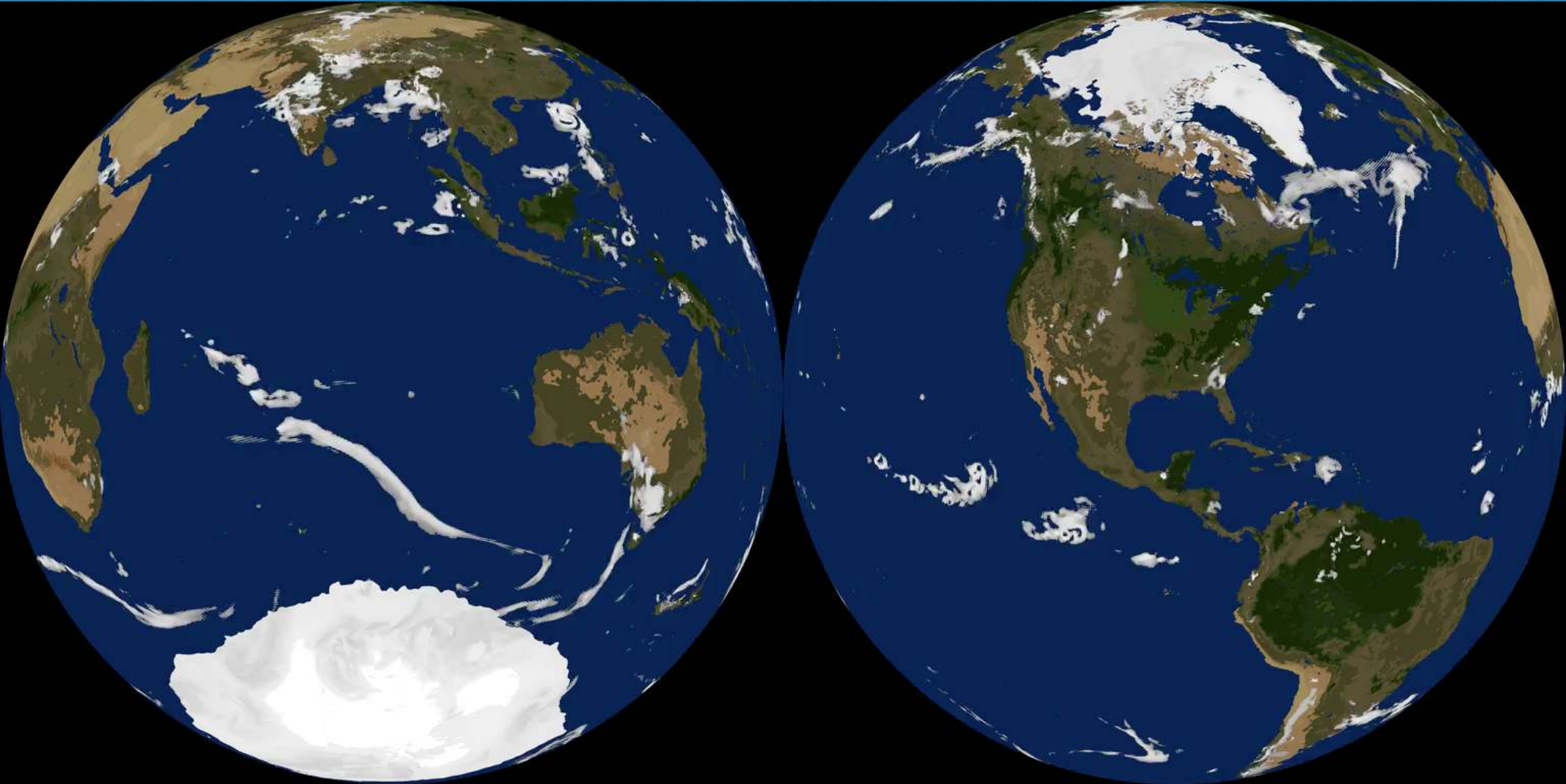


System for Atmospheric Modeling (SAM)

Grid spacing 4 km (0.039x0.039 deg)

Aug 1 - Sep 10 2016

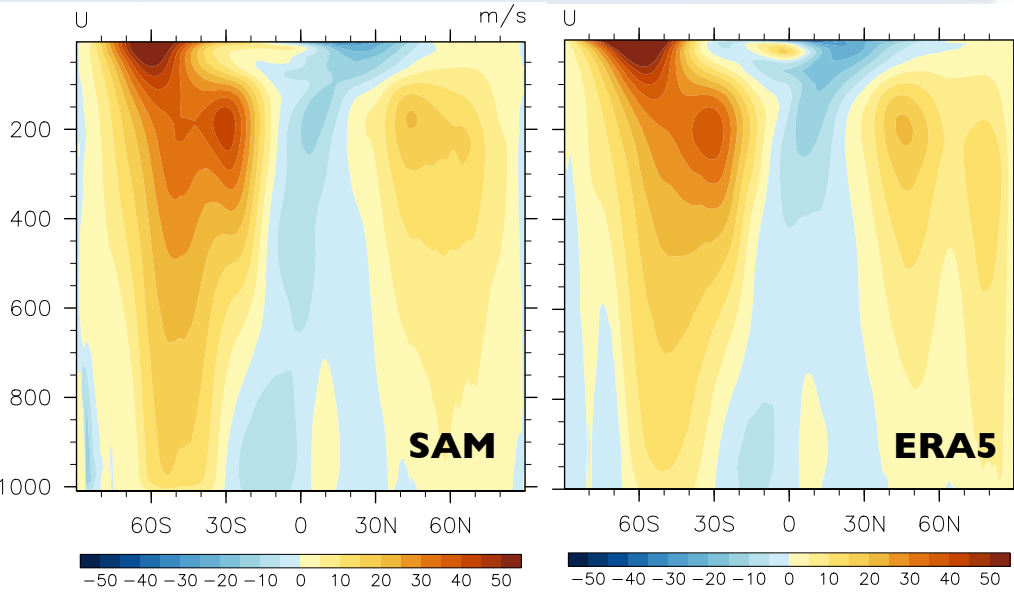
Cloud Albedo



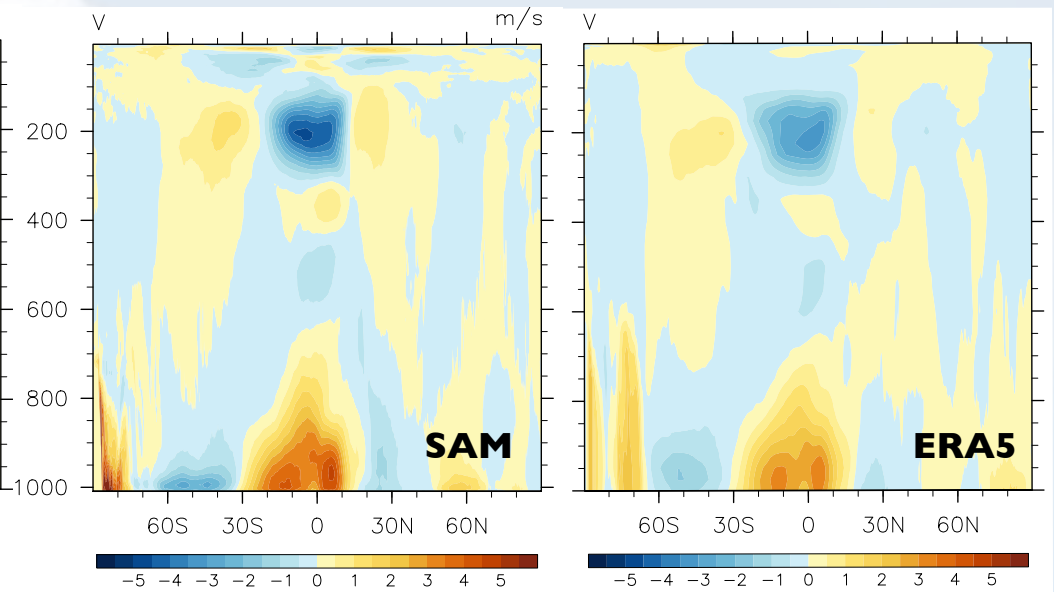
Performance: 6 times faster than real time on 4608 cores

Zonal-mean fields averaged over days 10-40

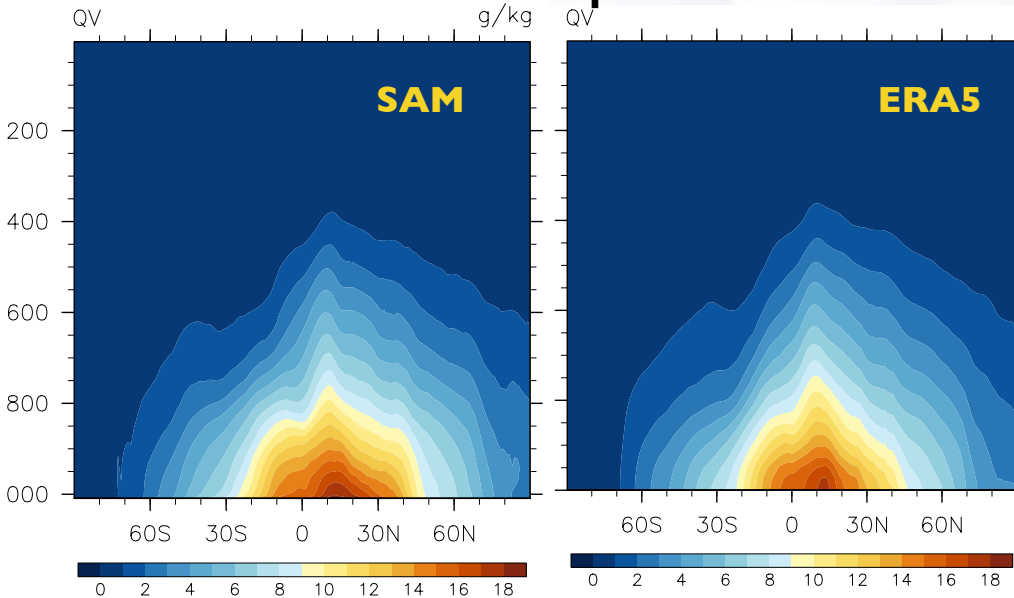
Zonal Wind



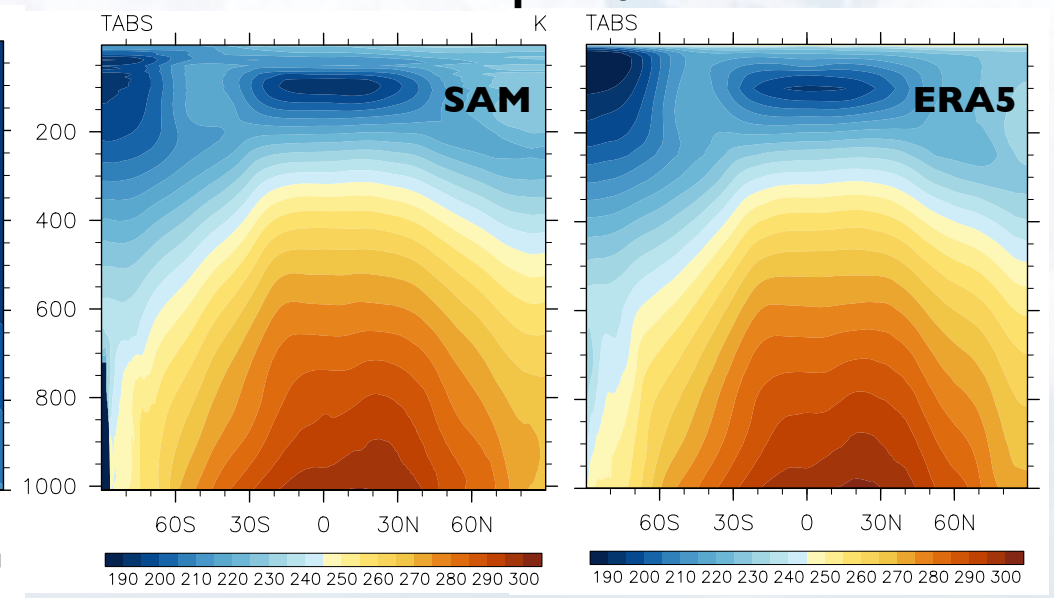
Meridional Wind



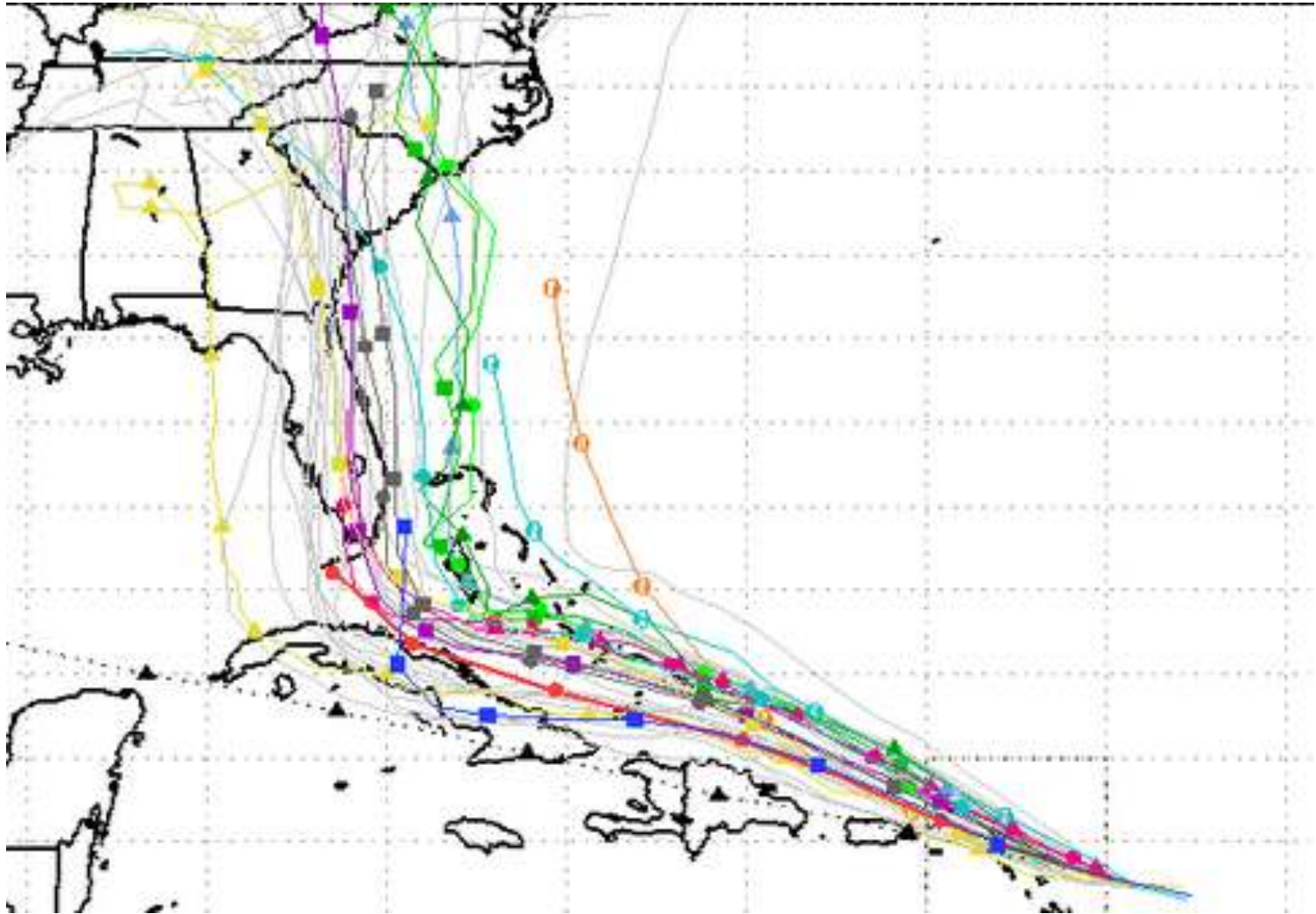
Water Vapor



Temperature



Hurricane Irma (2017)



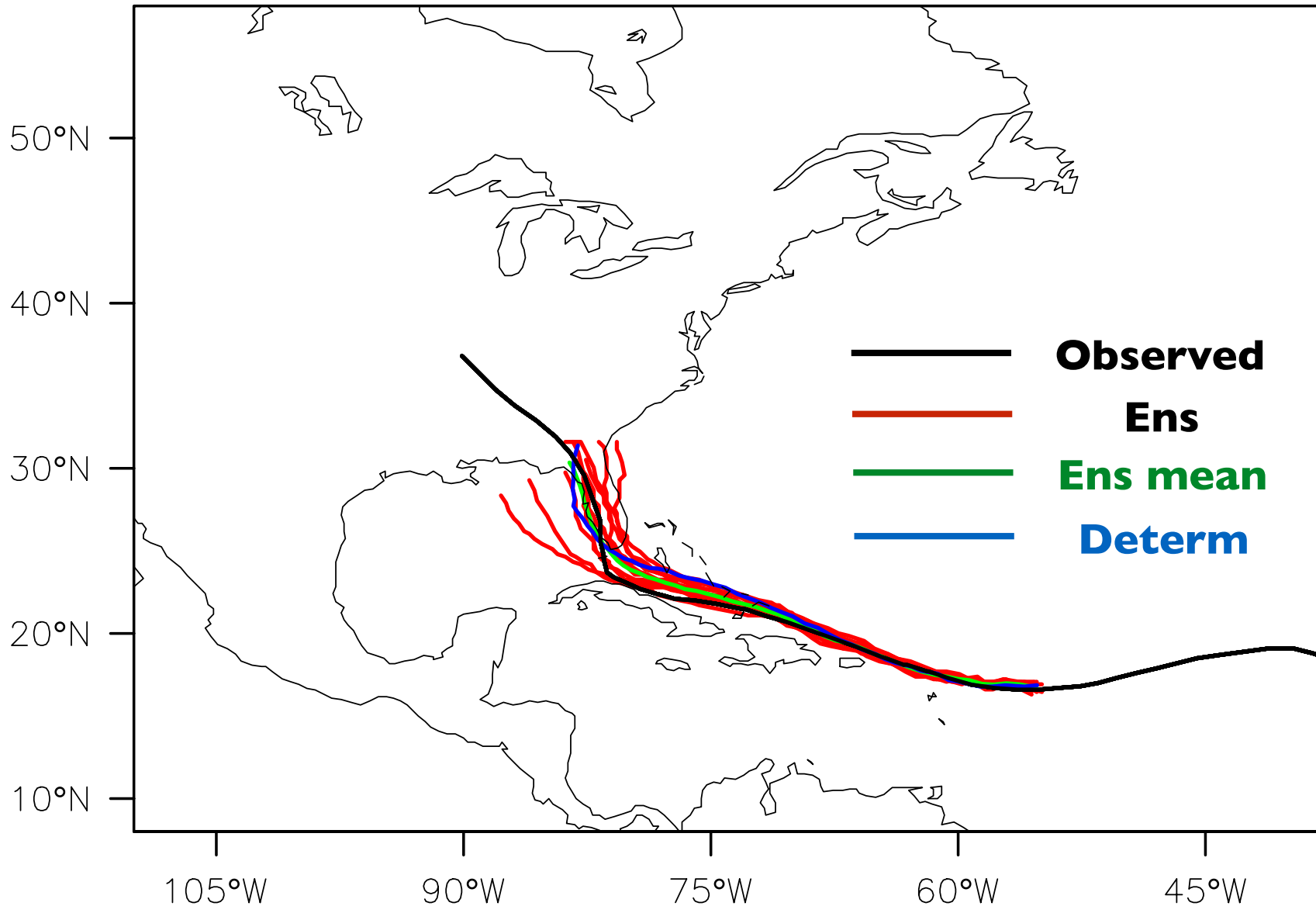
'Spaghetti' plot for forecast of Irma on Sep 5, 2017

SAM's 7-day forecast of hurricane Irma (2017)

Initialization: 10-member ensemble from ERA5 on Sep 5, 2017

Ensemble resolution: 17 km (50 times faster real time)

Deterministic forecast: 4 km (6 times faster real time)

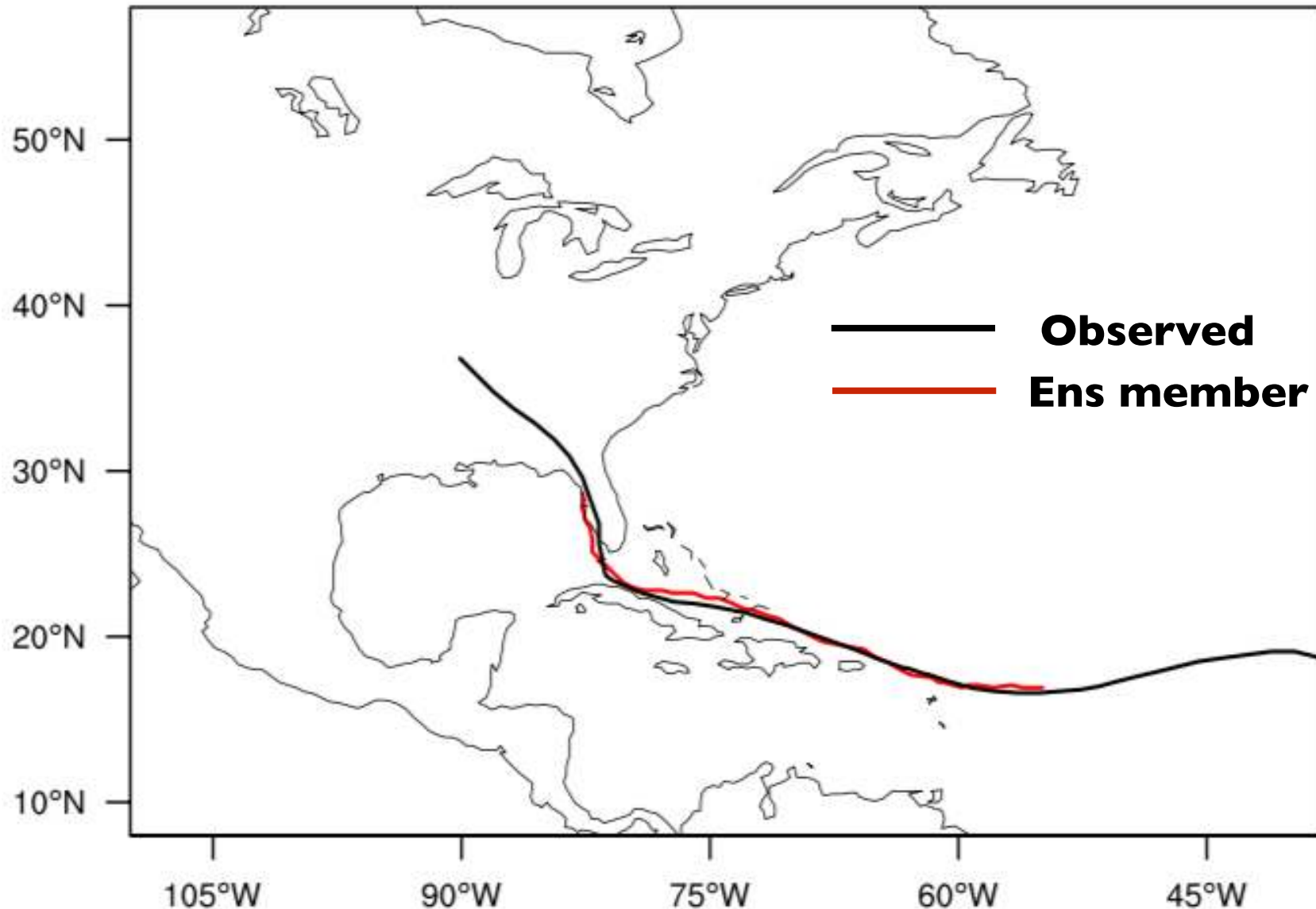


7-day forecast of hurricane Irma (2017)

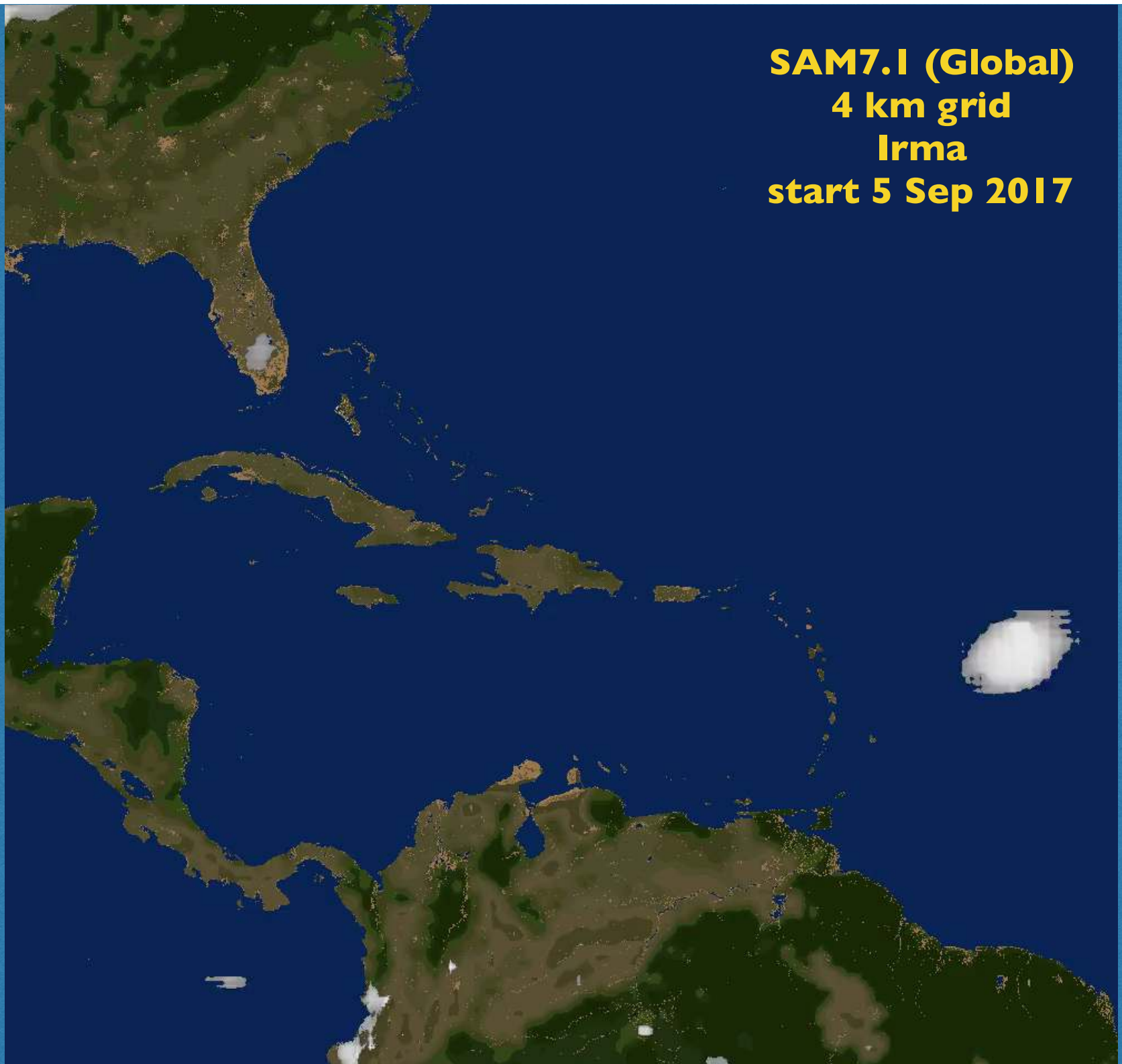
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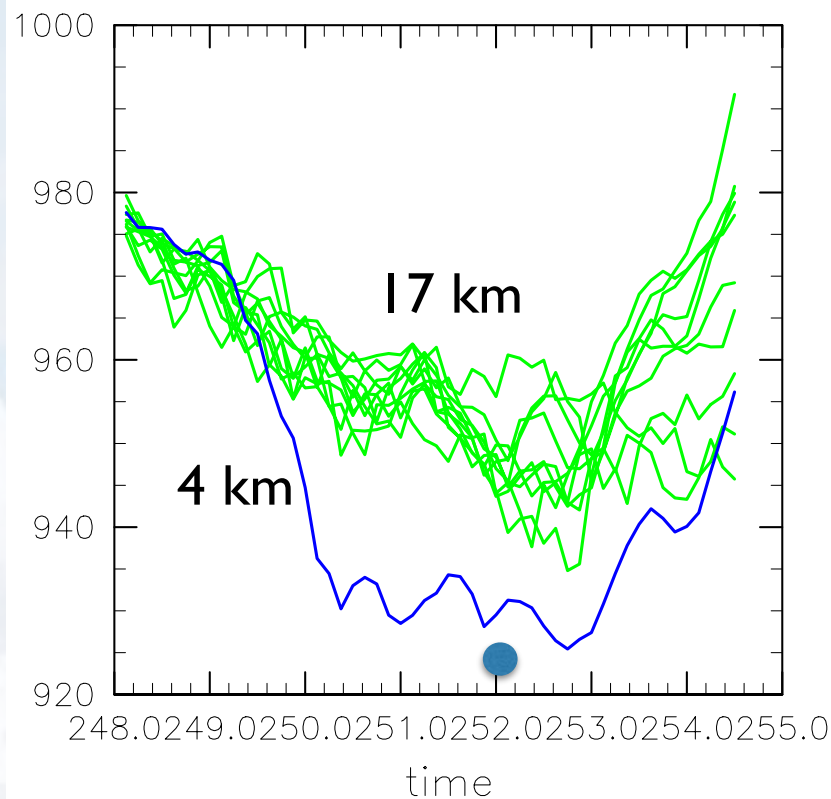


SAM7.1 (Global)
4 km grid
Irma
start 5 Sep 2017

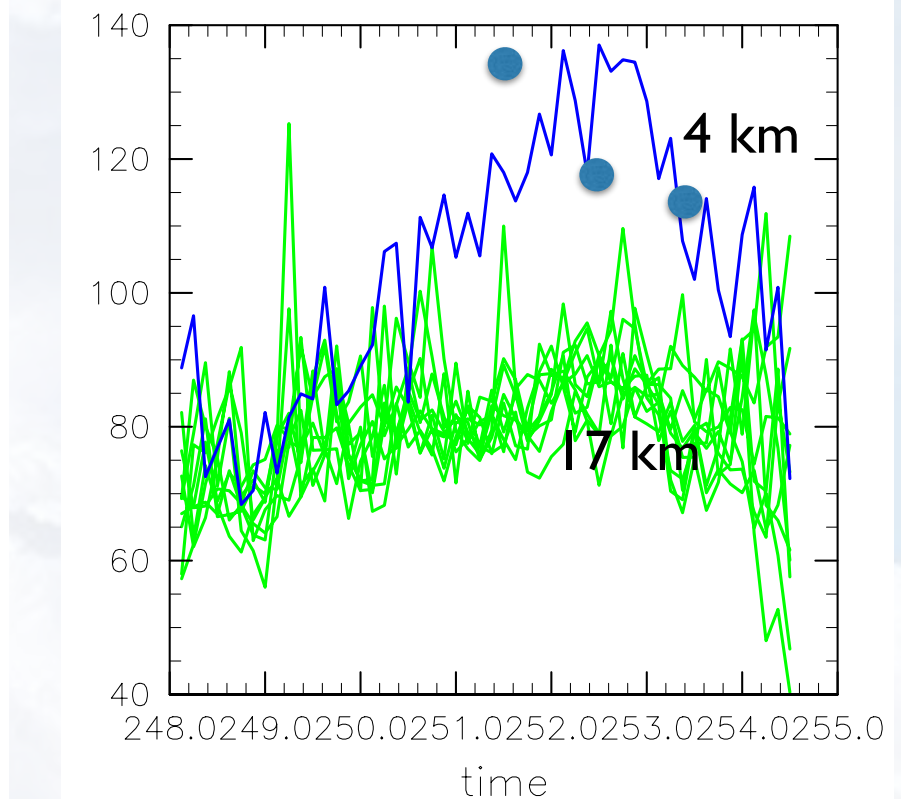


7-day forecast of hurricane Irma (2017)

Minimum Sfc Press, mb



Maximum Sfc Wind, kt

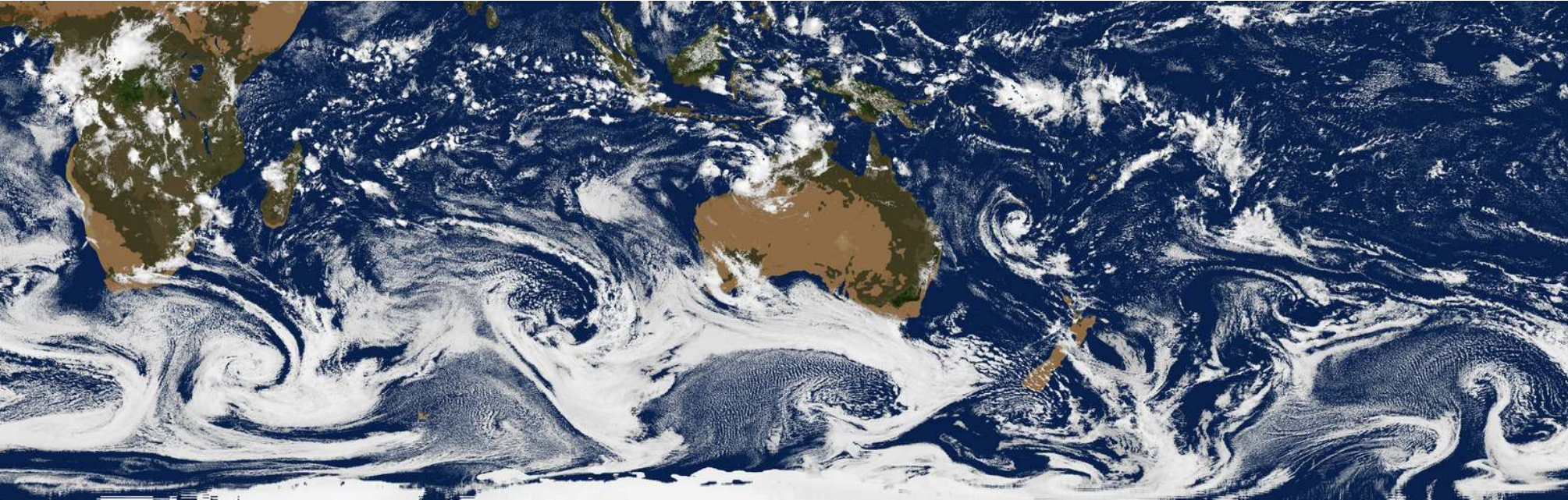


● Obs

**17-km grid simulation underestimate hurricane intensity.
4-km grid intensity and surface pressure agree well with
observations.**

I-day forecast of SOCRATES

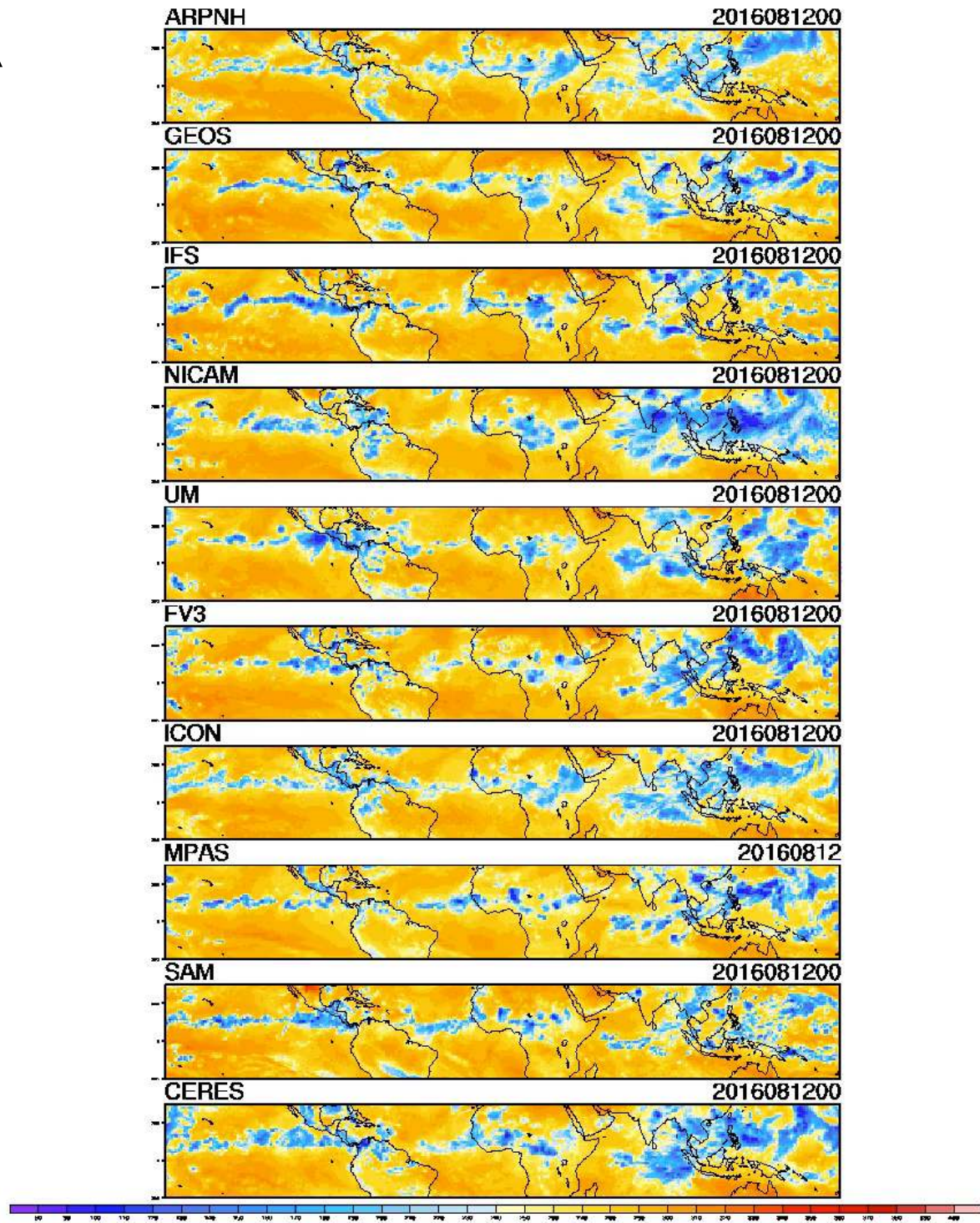
Feb 17 00Z



**Preliminary results of
Tropical-cluster Statistics
in DYAMOND Models**



1°x1° OLR

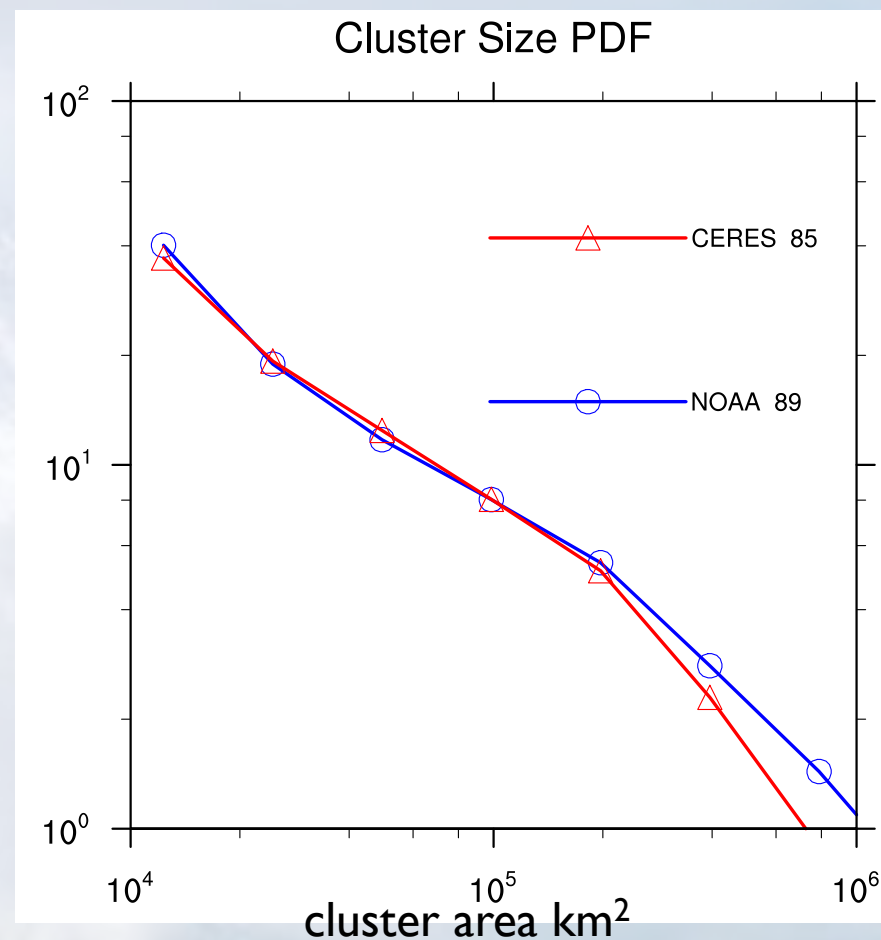


Data

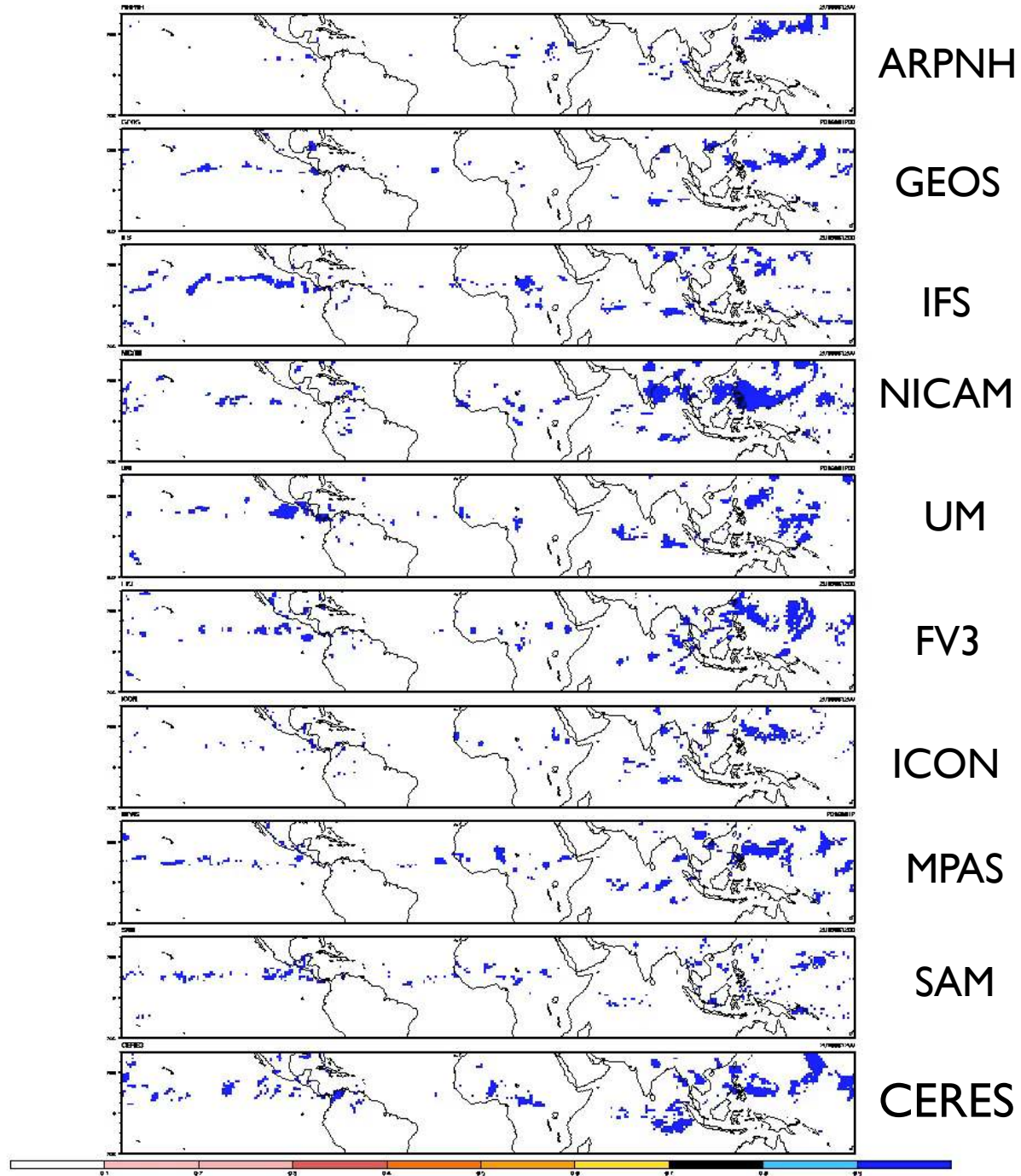
- **DYAMOND models**
 - **OLR, ASR, Prec, PW on uniform 0.1° grid**
- **There is no equivalent high-resolution (space and time) observed global OLR**
 - **NOAA $11\ \mu\text{m}$ 0.07° TBB, hourly**
 - **CERES 1° OLR hourly**
- **Which threshold for OLR to use for identifying deep convection?**
 - **400 mb temperature in Tropics is 254 K**
 - **Inoue et al (2008), for NICAM used $210\ \text{W/m}^2$ from cumulative histograms of OLR and TBB**
 - **This study: Use NOAA 0.07° TBB averaged to 1° and find CERES OLR threshold that produces similar PDF of cluster sizes**

Data

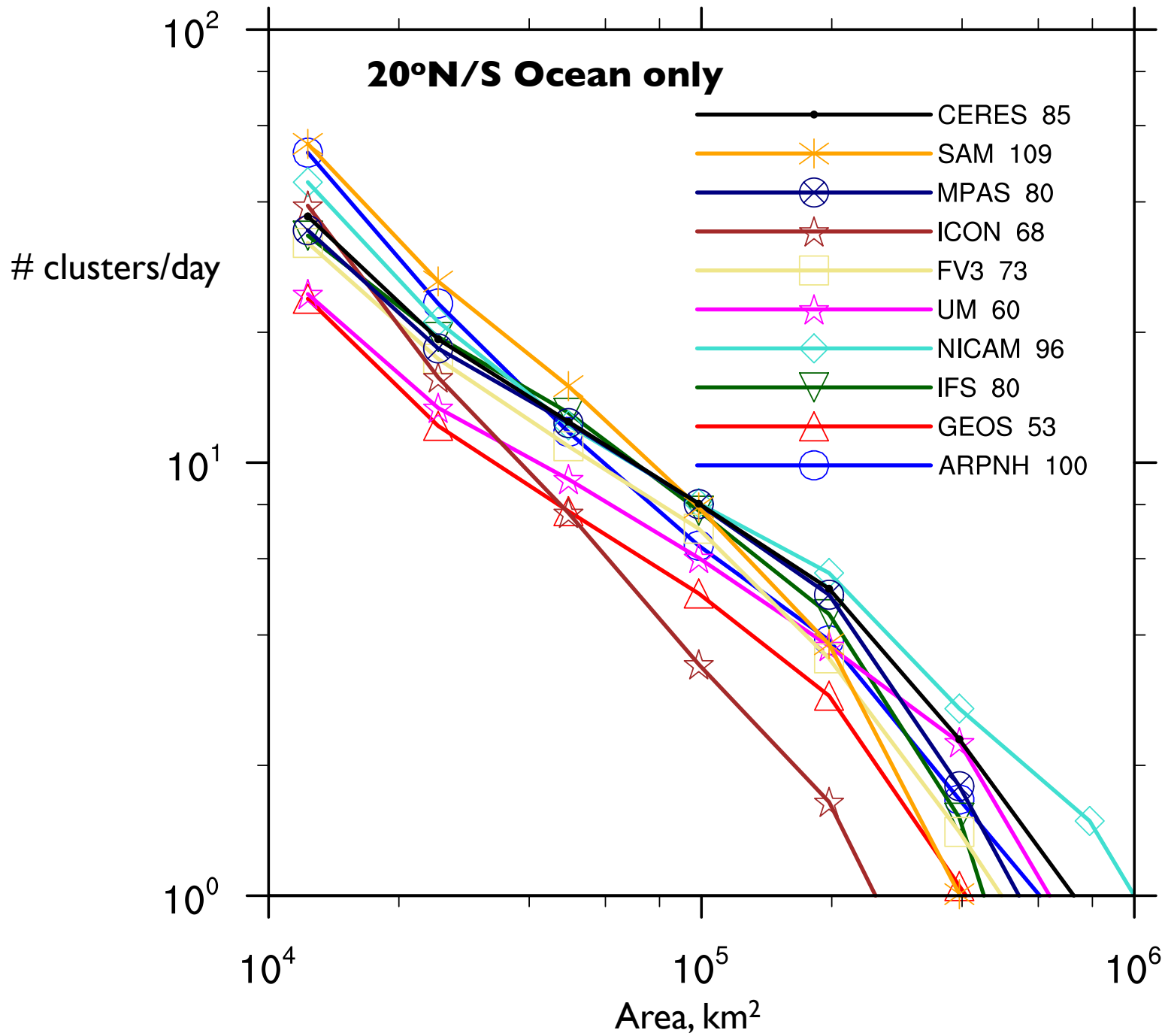
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 - **This study: Use NOAA 0.07° TBB averaged to 1° and find CERES OLR threshold that produces similar PDF of cluster sizes**
 - **Best correspondence for OLR = 160 W/m²**



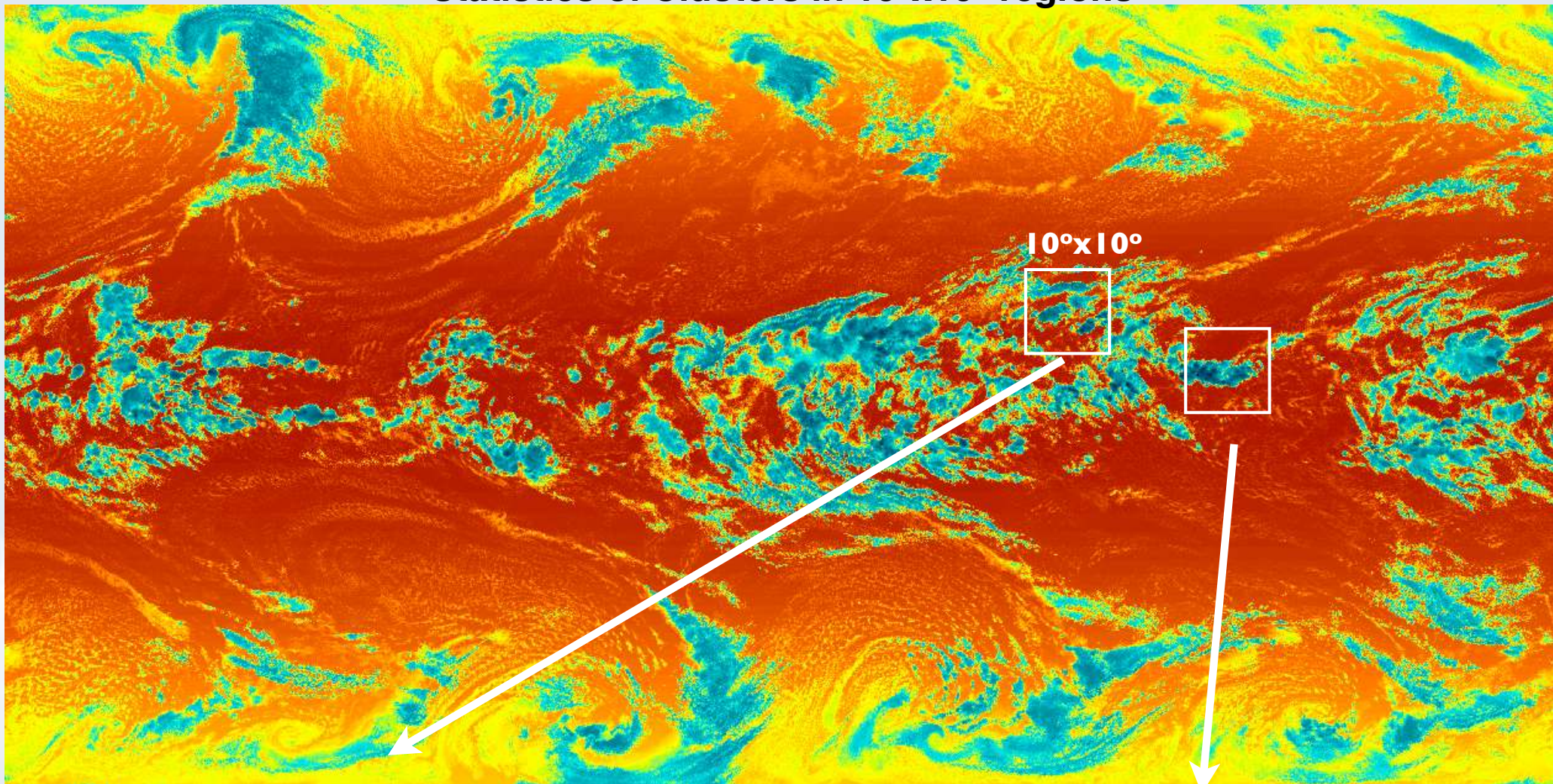
1° OLR < 160 W/m²



Cluster Size PDF

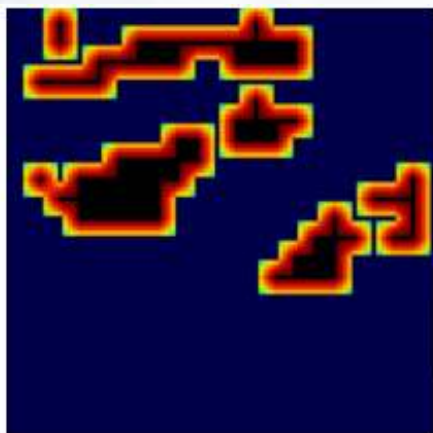


Statistics of Clusters in 10°x10° regions

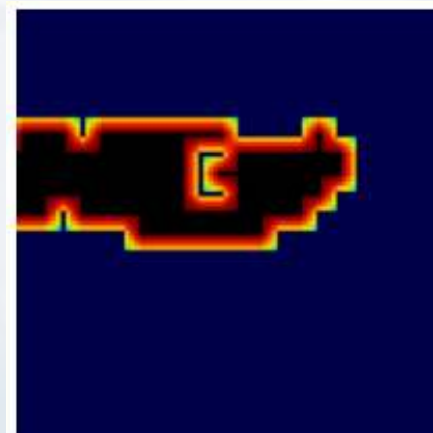


10°x10°

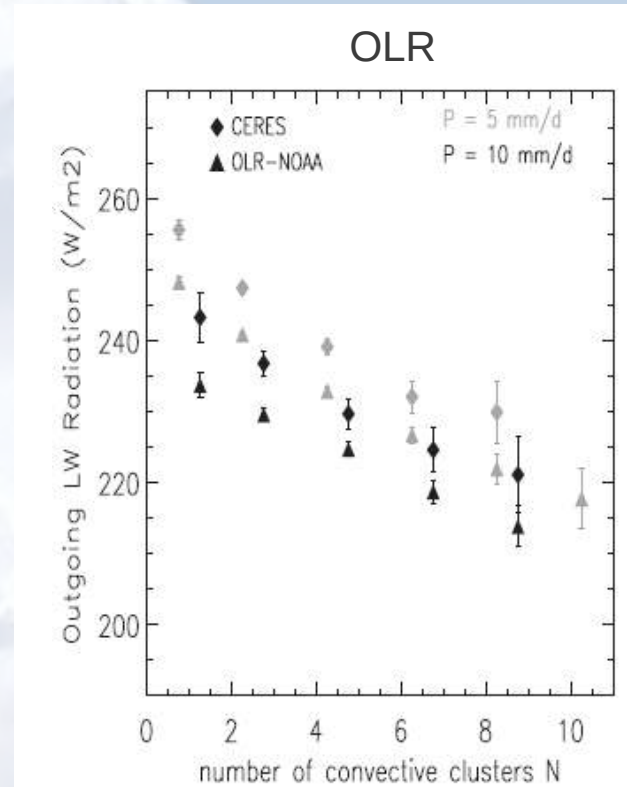
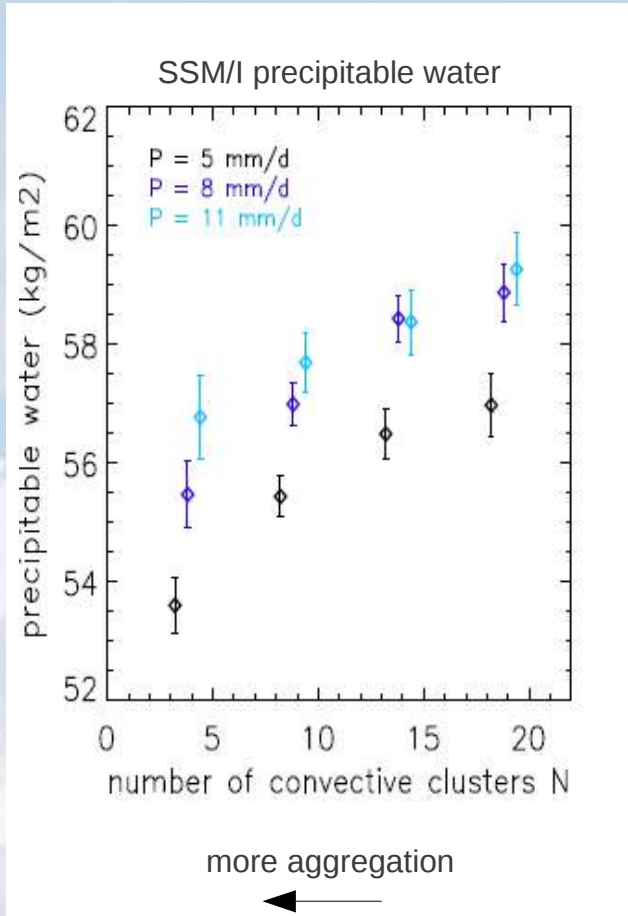
N=6



Less Aggregated More
Tb < 240K, 50-km resolution

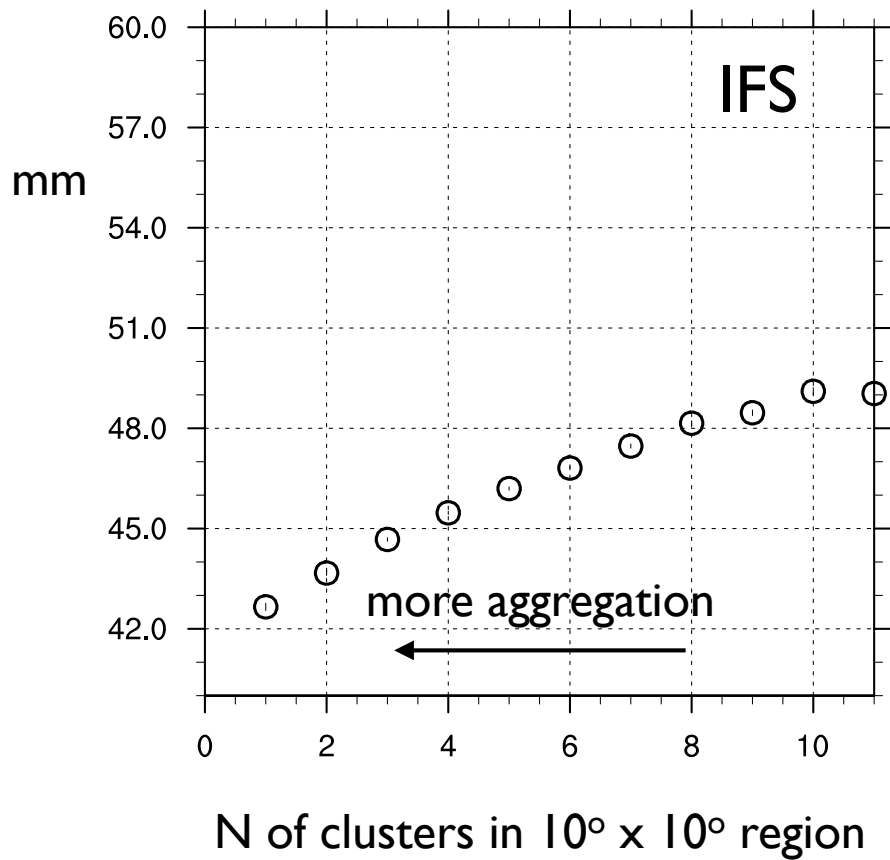


N=1

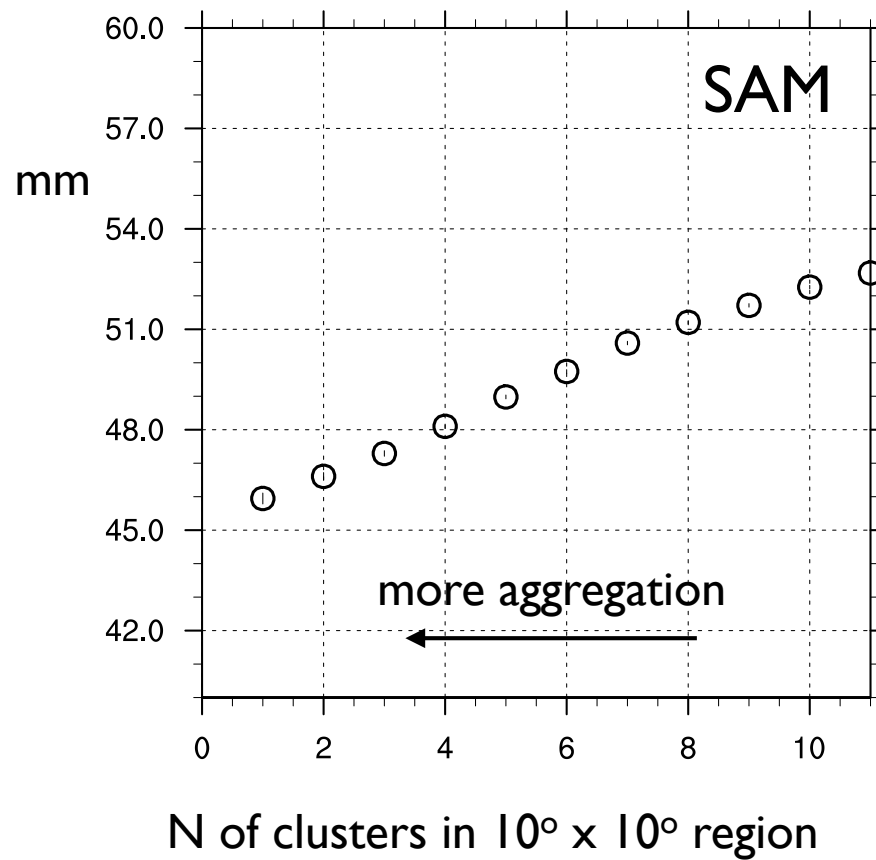


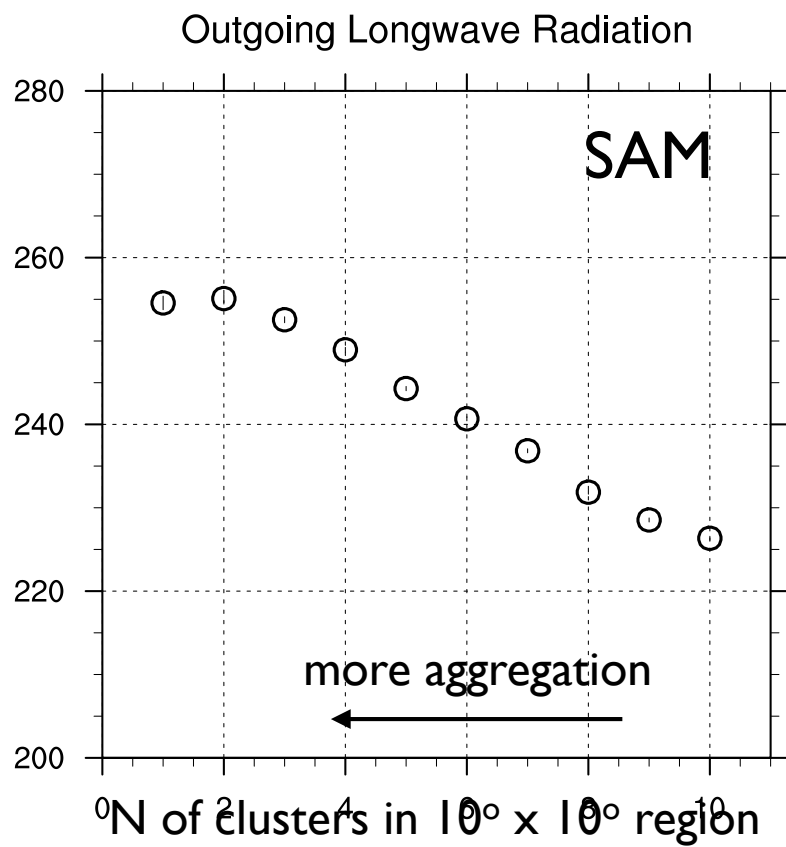
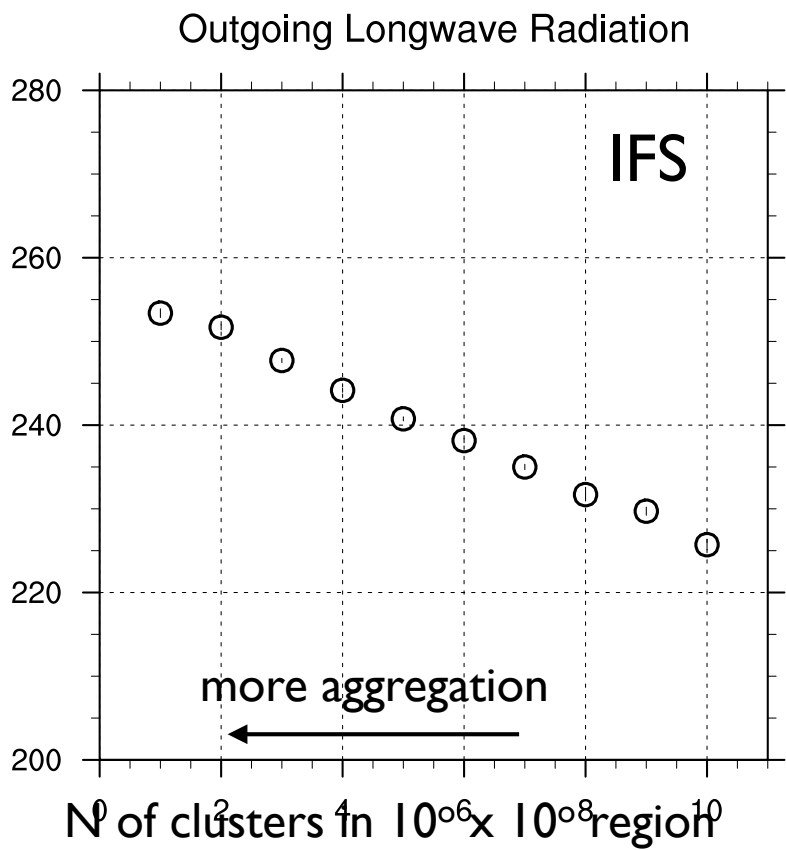
Tobin et al (2012, 2013):

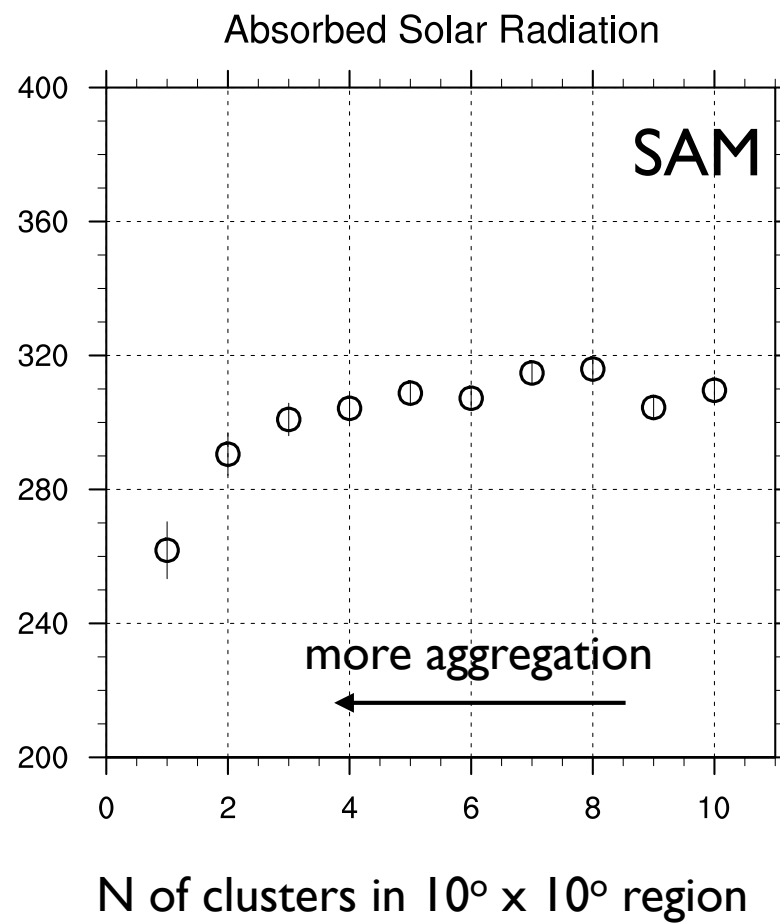
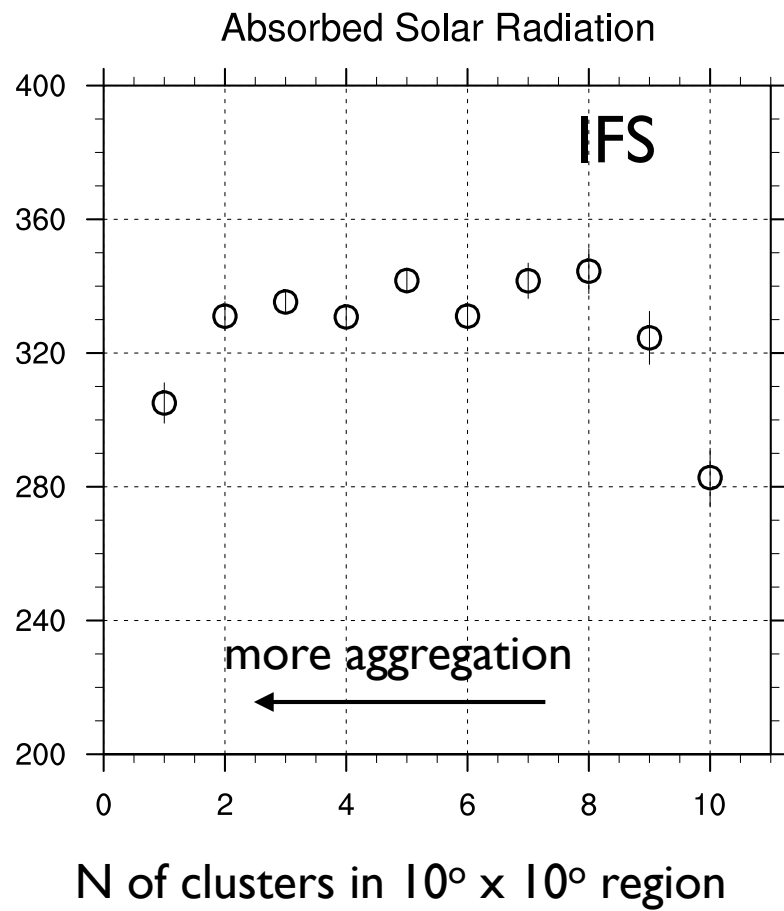
Column Vapor



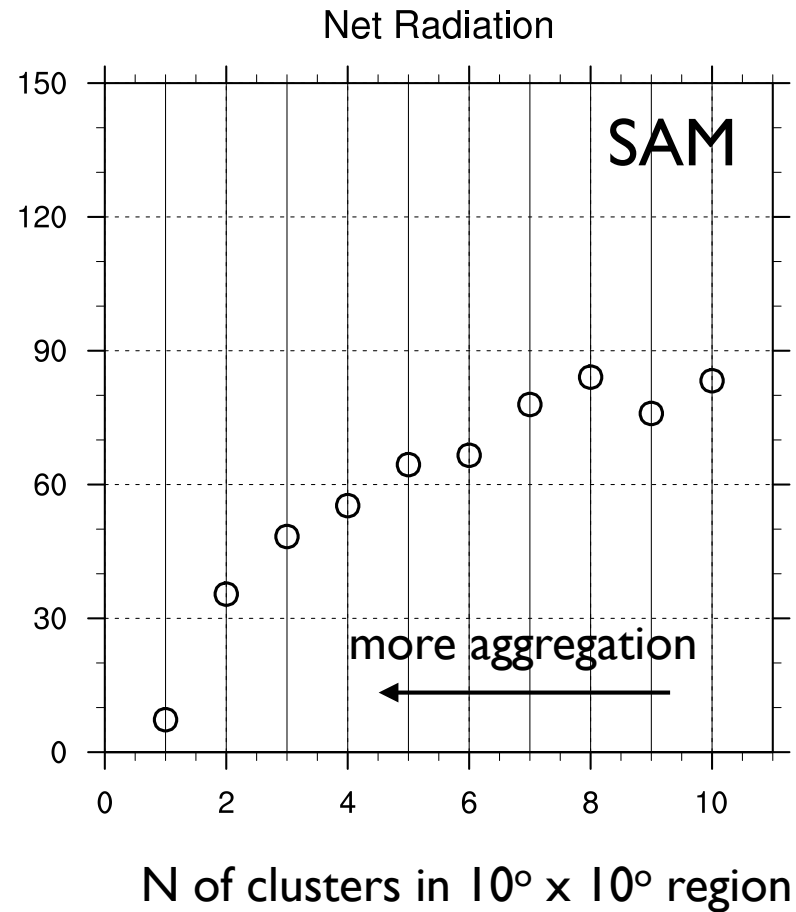
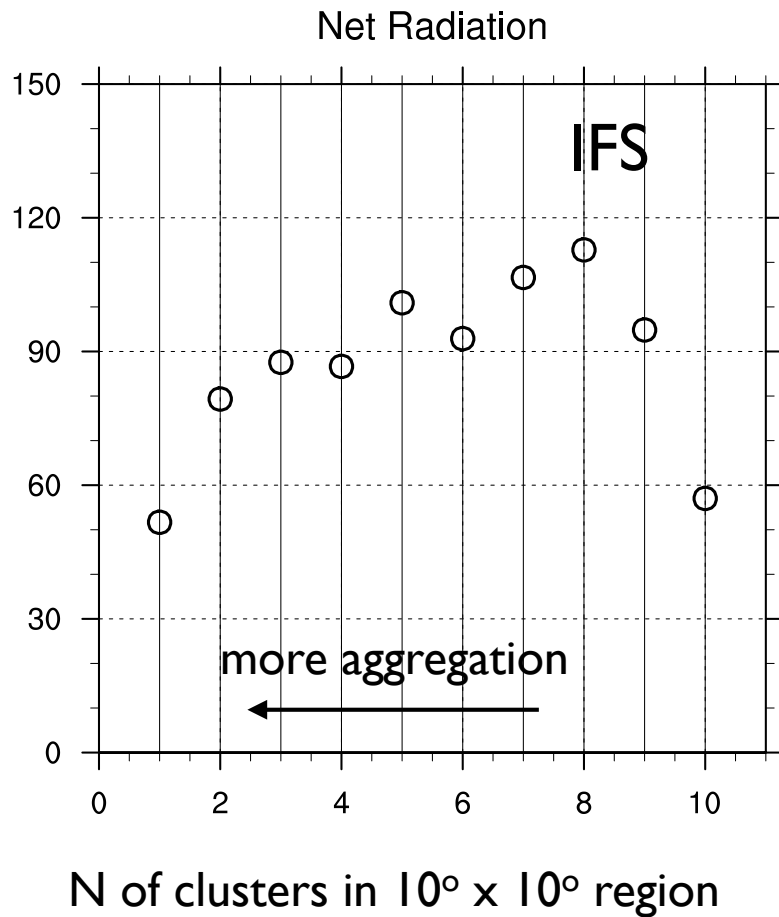
Column Vapor







ASR - OLR at TOA



Organized convection seems to cool the climate system?